## 1 The limits, capabilities, and potential for life detection with MinION sequencing in a

# 2 paleochannel Mars analogue

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### 24 Abstract

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26 No instrument capable of direct life detection has been included on a mission payload to 27 Mars since NASA's Viking missions in the 1970s. This prevents us from discovering if life is or 28 ever was present on Mars. DNA is an ideal target biosignature since it is unambiguous, non-29 specific, and readily detectable with nanopore sequencing. Here, we present a proof-of-concept 30 utilization of the Oxford Nanopore Technologies (ONT) MinION sequencer for direct life 31 detection and show how it can complement results from established space mission instruments. 32 We used nanopore sequencing data from the MinION to detect and characterize the microbial 33 life in a set of paleochannels near Hanksville, Utah, USA, with supporting data from X-ray 34 diffraction, reflectance spectroscopy, Raman spectroscopy, and Life Detector Chip (LDChip) 35 microarray immunoassay analyses. These paleochannels are analogues to Martian sinuous ridges. 36 The MinION-generated metagenomes reveal a rich microbial community dominated by Bacteria 37 and containing radioresistant, psychrophilic, and halophilic taxa. With spectral data and LDChip 38 immunoassays, these metagenomes were linked to the surrounding Mars analogue environment 39 and potential metabolisms (e.g. methane production and perchlorate reduction). This shows a 40 high degree of synergy between these techniques for detecting and characterizing biosignatures. 41 We also resolved a prospective lower limit of  $\sim 0.001$  ng of DNA required for successful 42 sequencing. This work represents the first determination of the MinION's DNA detection limits 43 beyond ONT recommendations and the first whole metagenome analysis of a sinuous ridge 44 analogue.

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## 47 **1. Introduction**

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49 The search for biosignatures is increasingly important in Mars exploration missions and 50 will be a primary objective for the upcoming ExoMars 2020 and Mars 2020 missions (McLennan 51 et al. 2012; Mustard et al. 2013a; Mustard et al. 2013b; Summons et al. 2014). In order to test 52 strategies, techniques, logistics, and instruments, terrestrial analogue sites are used to 53 approximate field operations and further understand the corresponding extraterrestrial body's 54 environment, geology, and potential for life and biosignature preservation (Thiel et al. 2011). 55 Hanksville, Utah, USA is situated near a number of diverse Mars analogues, including 56 segmented and inverted anastomosing paleochannels exhumed from the Late Jurassic Brushy 57 Basin Member of the Morrison Formation (Williams et al. 2007; Clarke and Stoker 2011; Ehrenfreund *et al.* 2011). Like other inverted fluvial channels, the Hanksville paleochannels have 58 59 resulted from relief inversion, in which materials in topographic lows are more resistant to 60 erosion than their surroundings (Williams et al. 2007; Williams et al. 2009; Clarke and Stoker 61 2011). Because fluvial channels are comprised of higher-permeability sediments than their 62 floodplain surroundings, the water flow can lead to cementation of the channels. Cemented 63 fluvial channels can then be exhumed by denudation of the surrounding landscape, resulting in 64 inversion of the former topography (Williams et al. 2009; Clarke and Stoker 2011). The Hanksville paleochannels were exhumed in this way, and the sedimentary sandstone of the 65 channels is capped with erosion-resistant coarse sandstones and fine-grained conglomerates 66 67 (Williams et al. 2007; Williams et al. 2009). This capping protects the underlying clay 68 floodplains and preserves the fluvial sediments, as well as sedimentary structures and 69 biosignatures (Williams et al. 2007; Clarke and Stoker 2011). The Hanksville paleochannels are

70 strong geomorphological analogues to sinuous ridges on Mars, which is a non-specific term for 71 elongated ridge structures on the Martian surface. They may have formed from inverted stream 72 channels but could also result from other, non-hydrological processes (e.g. wrinkle ridges) 73 (Williams et al. 2009; Williams et al. 2013). 74 Dating from the Noachian Period to the Amazonian Period on Mars, sinuous ridges 75 represent compelling evidence for a history of prolific water flow (Clarke and Stoker 2011; Balme et al. 2015; Davis et al. 2016). They are distributed globally, but are particularly abundant 76 77 in Arabia Terra, the northernmost region of Mars' heavily cratered southern highlands (Williams 78 et al. 2009; Clarke and Stoker 2011; Davis et al. 2016). One prominent sinuous ridge in this area 79 is Aram Dorsum, a former candidate landing site for ExoMars 2020 (downselected in March 80 2017). In November 2016, the Hanksville paleochannels were chosen for CanMars 2016, a 81 simulation Mars sample return mission led jointly by the Canadian Space Agency and Western 82 University (Osinski et al. 2018). Mars sample return will take its first steps with the Mars 2020 83 rover, which will identify and cache organic samples and water-formed samples for potential 84 return to Earth in the 2030s (Mustard et al. 2013a; Mustard et al. 2013b; Summons et al. 2014; 85 Beegle *et al.* 2015). The National Research Council (USA) and the Mars 2020 Science Definition Team both 86

state that the determination if life is or ever was present on Mars is the prime focus of the 2013 –
2022 decade (NRC 2012; Mustard *et al.* 2013b). Yet neither ExoMars 2020 nor Mars 2020
contain instrumentation capable of detecting definitive biosignatures, which are substances,
objects, or patterns that could not form via abiotic processes (Mustard *et al.* 2013b; Hays *et al.*2017; Sephton 2018). Nucleic acids, including DNA, are strong definitive biosignatures given
the virtual impossibility of their generation in the absence of life (Mustard *et al.* 2013b; Neveu *et*

93 al. 2018). In missions focusing on Mars, DNA is an especially good target. There is some 94 evidence for the exchange of possible microbe-carrying rocks between Earth and Mars during 95 the Late Heavy Bombardment period ~ 4 billion years ago, when both planets shared a warm and 96 wet environment capable of harbouring life (Mileikowsky et al. 2000). Overall geological and 97 chemical similarities between Earth and Mars indicate that any potential Martian life could share 98 fundamental biochemistry with terrestrial life because the emergence of inorganic-based biology 99 would require conditions and chemistries vastly different from the two rocky planets (e.g. gas 100 giants and icy moons) (Mustard et al. 2013b). There is also a widespread occurrence of organic 101 molecules in interstellar space (Pace 2001; Meinert et al. 2016) and putative hydrogen-rich 102 aliphatics have recently been detected in Mars' Gale Crater (Eigenbrode et al. 2018). While it is 103 certainly not definite that life on Mars (or anywhere else in our solar system) would share its 104 biochemical phenotype with Earth, we should include the capability to detect these features on 105 Mars exploration missions. We must search for what we know before we can rule out its 106 impossibility. This is especially true with increasing activity in both the private and public space 107 sectors that could compromise future life detection and science activities on Mars (Fairén et al. 108 2017; Musk 2017; Spry et al. 2018).

As well as detecting DNA, life detection instrumentation should have the capacity for miniaturization and low power requirements to ease deployment on a rover or lander vehicle. It should also have proven functionality in terrestrial analogue environments (Ellery *et al.* 2002). The Oxford Nanopore Technologies' (ONT) MinION is a promising candidate for life detection based on these requirements. The MinION is a miniaturized device designed to sequence DNA, RNA, and potentially proteins (Loman and Watson 2015; Goordial *et al.* 2017; Leggett and Clark 2017). The MinION uses nanopore technology to generate sequences (Loman and Watson

116 2015; Leggett and Clark 2017). It is portable and can operate in difficult environmental 117 conditions, including the high Arctic, Antarctic, tropical rainforests, and the microgravity of the 118 International Space Station (Edwards et al. 2016; John et al. 2016; Castro-Wallace et al. 2017; 119 Goordial et al. 2017; Johnson et al. 2017; Menegon et al. 2017). Indeed, the MinION was the 120 first sequencer to operate in space as part of the Biomolecule Sequencer project in 2016 (John et 121 al. 2016; Castro-Wallace et al. 2017), making it a promising candidate for direct life detection in 122 space missions. Its small size would make it easier and cheaper to develop than larger payloads, 123 it can function in a controlled environment like the heated interior chamber of a rover, and it can 124 directly detect unambiguous signs of life through nucleic acid sequencing (Carr et al. 2017; 125 Goordial et al. 2017). This would enable the best selection of samples for future caching and 126 potential return to Earth (Goordial et al. 2017). If DNA is detected on a mission, the ability of 127 the MinION to generate sequences in real time also means that contamination from terrestrial 128 microorganisms can be quickly ruled out. However, ONT requires a minimum input of 400 ng of 129 DNA without PCR amplification and 1 ng of DNA with PCR amplification for MinION 130 sequencing. The low biomass present in many Mars analogues, like high elevation University 131 Valley permafrost ( $10^3$  cells/g) (Goordial *et al.* 2016) and Lost Hammer Spring sediment ( $10^5$ 132 cells/g) (Niederberger *et al.* 2010), corresponds to just ~ 0.01 - 1 ng of DNA based on an 133 average bacterial genome size of 2-5 Mbp and  $\sim 2.5-5$  fg of DNA per genome (Dolezal *et al.* 2003; Raes et al. 2007; Goordial et al. 2017), meaning that DNA detection limits of MinION 134 135 sequencing need to be explored beyond ONT recommendations. 136 We tested the life detection capability of the MinION with samples from a high fidelity 137 analogue environment, the Hanksville paleochannels. We determined DNA detection limits

using a high biomass endolith sample and generated a metagenomic profile of this site with

139	MinION sequences. Simultaneous analyses with an X-ray diffractometer, micro-reflectance
140	spectrometer, and micro-Raman spectrometer and protein detection with the LDChip (Life
141	Detector Chip) (Parro et al. 2005) were also performed to complement MinION results and to
142	establish baseline data for a paleochannel environment. These instruments were chosen because
143	they are analogues to current and future payload instruments. An X-ray diffractometer forms the
144	core of the Curiosity rover's CheMin instrument, a micro-reflectance spectrometer acts as an
145	approximation for ExoMars 2020's Mars Multispectral Imager for Subsurface Studies
146	(Ma_MISS), and a micro-Raman spectrometer approximates ExoMars 2020's Raman Laser
147	Spectrometer (RLS) (Rull et al. 2017) and Mars 2020's SuperCam and SHERLOC (Scanning
148	Habitable Environments with Raman and Luminescence for Organics and Chemicals). The
149	LDChip approximates the Life Marker Chip (LMC), a former Pasteur payload candidate for
150	ExoMars 2020. These instruments characterize other biosignatures (e.g. isotopes, organics, and
151	polymeric biomarkers), contextualize the environment and geologic setting, and help reveal
152	metabolic potential of the microbial community.
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154	2. Materials and Methods
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156	2.1. Sample site and collection
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158	The study site is located approximately 7.8 km NW of Hanksville, Utah, USA
159	(38.41709915, -110.78539335; Fig. 1). It is part of the Colorado Plateau and contains a diverse
160	array of early Jurassic to late Cretaceous-aged evaporitic and fluvial sediments. The geology of
161	the site is dominated by quartz, calcite, and gypsum, and has been shaped by fluvial and aeolian

162	processes (Martins et al. 2011; Stoker et al. 2011). It is an arid desert with an average annual
163	temperature of 11.9 °C and 144 mm of precipitation (Martins et al. 2011). Samples were
164	collected to represent a range of geomorphology and biomasses (e.g. desiccated regolith, high
165	biomass endoliths). Regolith was aseptically collected with a hand shovel and harder rocks were
166	broken off of larger concretions with a rock hammer. All collection tools were sterilized with 70
167	% ethanol and a lighter. Latex gloves were worn during collection and samples were loaded into
168	sterile Whirl Pak plastic bags. Samples were transported from the Utah site to Montreal, Quebec,
169	Canada at ~ 5 $^{\circ}$ C and subsequently stored at -30 $^{\circ}$ C while waiting for analysis at McGill
170	University (MinION sequencing), the University of Winnipeg (X-ray diffraction, reflectance
171	spectroscopy, Raman spectroscopy), and the Centro de Astrobiología (INTA-CSIC) (LDChip
172	immunoassays).
173	
174	2.2. DNA extraction and MinION sequencing
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176	DNA was extracted from all samples with the PowerLyzer PowerSoil DNA Isolation Kit
177	(Mo Bio Laboratories, Inc., Carlsbad, CA) according to the manufacturer's protocol and eluted
178	in 50 $\mu$ l of nuclease-free water. Extractions were stored at -20 °C until sequencing. Extractions
179	were prepared for 1D single strand sequencing using the Rapid Low Input by PCR Barcoding kit
180	(SQK-RLB001) and all samples were sequenced on 1 R9.5 FLO-MIN107 flow cell during 1 run
181	(Table 1).
182	DNA detection limits were established using endolith sample E4, since it had the highest
183	percent of sequences classified in the Joint Genome Institute's Integrated Microbial Genomes
184	and Microbiomes Expert Review (JGI IMG/M ER) system (Markowitz et al. 2011) and the

185	second highest number of sequences produced (Table 2). In order to determine if the MinION
186	could distinguish false positives from true positives, a run containing no DNA (3 $\mu$ l of sterile
187	nuclease-free water) was also sequenced using the same method (SQK-RLB001 preparation) on
188	an R9.5 flow cell. The detection limit trial samples were prepared for sequencing using the Rapid
189	PCR Barcoding kit (SQK-RPB004), since the Rapid Low Input by PCR Barcoding kit (SQK-
190	RLB001) was discontinued soon after our metagenome analyses. Each trial was sequenced on a
191	separate R9.4 FLO-MIN106 flow cell; R9.5 chemistry was no longer recommended for 1D
192	single strand sequencing when we conducted our detection limit trials.
193	All sequences were basecalled with MinKNOW v1.7.7 and underwent the following
194	downstream processing: basecalled sequences were demultiplexed and trimmed using Porechop
195	version 0.2.3 (https://github.com/rrwick/Porechop) with parameters "barcode_threshold 60"
196	and "barcode_diff 1." Demultiplexed sequences were corrected and assembled with Canu
197	version 1.7 +75 (Koren et al. 2017) with parameters "corOutCoverage=10000,"
198	"corMhapSensitivity=high," "correctedErrorRate=0.16," and an assumed genome size of 4.7
199	Mbp (Raes et al. 2007). These sequences were then uploaded to JGI IMG/M ER (Markowitz et
200	al. 2011) for annotation. All datasets are public, and the JGI IMG/M ER submission IDs are
201	included in Table 1 for the shotgun metagenome and in Table 2 for the detection limit sequences.
202	A bubble plot of the whole shotgun metagenome was constructed using the following script:
203	https://github.com/alex-bagnoud/OTU-table-to-bubble-plot.git.
204	We also performed 16S rRNA gene sequencing with an Illumina MiSeq according to the
205	protocol described in Kozich et al. (2013). All sequences were classified using the mothur
206	software package (http://www.mothur.org/wiki/MiSeq_SOP) and the SILVA v132 reference
207	files. Sequence data is publicly available on NCBI with the following BioProject accession

208	number: PRJNA544288. OTU tables and taxonomy classification were used to construct bubble
209	plots with the following script: https://github.com/alex-bagnoud/OTU-table-to-bubble-plot.git.
210	Principal coordinates analysis (PCoA), PERMANOVA values, and PHRED quality score plots
211	were constructed in CLC Genomics Workbench v12.0 (https://www.qiagenbioinformatics.com/).
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213	2.3. Powder X-ray diffraction (XRD)
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215	X-ray diffraction (XRD) data was acquired in continuous scans from 5 to 80° 2 $\theta$ on a
216	Bruker D8 Advance with a DAVINCI automated powder diffractometer from crushed aliquots of
217	sample. It uses a Bragg-Brentano goniometer with a theta-theta system. The setup was equipped
218	with a 2.5° incident Soller slit, 1.0 mm divergence slit, a 2.0 mm scatter slit, a 0.6 mm receiving
219	slit, and a curved secondary graphite monochromator. Diffracted X-rays were collected by a
220	scintillation counter at increments of 0.02° and an integration time of 1 second per step. The line
221	focus Co X-ray tube was operated at 40 kV and 40 mA, using a take-off angle of 6°. Diffraction
222	patterns were interpreted using Bruker Diffracsuite EVA software and the International Center
223	for Diffraction Data Powder Diffraction File (ICDD-PDF-2) database.
224	
225	2.4. Reflectance Spectroscopy (UV-VIS-NIR, 350-2500 nm)
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227	Spectra of regolith samples were collected in a powdered and whole rock aliquot, and
228	multiple spots were analyzed on the endolith-bearing samples (both fresh and weathered rock,
229	and endolith surfaces). Long wave ultraviolet, visible and near-infrared (350-2500 nm)
230	reflectance spectra were measured with an Analytical Spectral Devices FieldSpec Pro HR

231	spectrometer ( $\Delta$ SD). This instrument has a spectral resolution between 2 and 7 nm (internally
231	spectrometer (ASD). This instrument has a spectral resolution between 2 and 7 mil (internally
232	resampled by the instrument to 1 nm). Spectra were collected at a viewing geometry of $i = 30^{\circ}$
233	and $e = 0^{\circ}$ with incident light being provided by an in-house 150 W quartz-tungsten-halogen
234	collimated light source. Samples were measured relative to a Spectralon® 100 % diffuse
235	reflectance standard and corrected for minor (< 2 %) irregularities in its absolute reflectance. We
236	acquired and averaged 200 spectra of the dark current, standard, and sample to provide sufficient
237	signal-to-noise for subsequent interpretation.
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239	2.5. Raman Spectroscopy (532 nm)
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241	Raman spectra were collected from multiple spots on both whole rock and powdered (<
242	$45\mu m$ ) samples (same as the reflectance spectra) using a B&WTek i-Raman-532-S instrument in
243	the Raman shift range of 175-4000 cm <sup>-1</sup> . This was done with a spectral resolution of ~ 4 cm <sup>-1</sup> at
244	614 nm with a 532 nm excitation energy provided by a $\sim$ 50 mW solid state diode laser. Raman-
245	scattered light was detected by a Glacier <sup>TM</sup> T, a high spectral resolution (0.08 nm)
246	thermoelectrically cooled (14 $^{\circ}$ C) CCD detector. The automatic integration time function (which
247	increases integration time incrementally, until the response is close to saturation) was used,
248	yielding an optimal signal-to-noise ratio (SNR). Measurements for each sample were made by
249	first acquiring a dark current spectrum, followed by measurement of the sample. Both
250	measurements were made using an identical viewing geometry, integration time, and number of
251	averaged spectra. Raman-shift calibration was monitored through regular measurements of a
252	polystyrene standard. The RUFF database (http://rruff.info/) was used for peak identification
253	(Downs 2006).

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### 2.6. LDChip (Life Detector Chip) Biomolecule Detection

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257 The LDChip (Life Detector Chip) is the core of the SOLID (Signs of Life Detector) 258 instrument (Parro et al. 2011) and consists of an antibody microarray composed of about 200 259 polyclonal antibodies (purified IgG fraction) designed to search for microbial molecular 260 biomarkers for environmental and space applications (Rivas et al. 2008). The LDChip contains 261 antibodies to: i) prokaryotic strains to cover the most abundant phylogenetic phyla and groups; 262 ii) proteins and peptides from key metabolic processes operating in extreme and "normal" 263 environments (e.g. nitrogen fixation, sulfate reduction/oxidation, iron reduction/oxidation, 264 methanogenesis); iii) universal microbial markers like peptidoglycan; and iv) crude extracts from 265 extreme environments (Rivas et al. 2008; Parro et al. 2011). The targets for the antibodies of this 266 study are described in (Sánchez-García *et al.* 2018) (personal communication). For this work, 267 LDChip microarrays were constructed, processed, and analyzed as detailed described in Blanco 268 et al. (2015, 2017). Briefly, samples (IgG fraction of each antibody) were printed in a triplicate 269 spot-pattern on the surface of epoxy-activated glass slides, fluorescently labeled, titrated, and 270 used in a mixture that consisted of 181 antibodies to reveal immunoreactions. 0.5 g of each 271 sample were resuspended into 2 mL of TBSTRR buffer (0.4 M Tris-HCl pH 8, 0.3 M NaCl, 0.1 272 % Tween 20) and used as a multianalyte-containing sample for fluorescent sandwich microarray 273 immunoassays (FSMI). LDChip images were analyzed and quantified by GenePix Pro Software 274 (Molecular Devices, CA, USA). The final fluorescence intensity (F) was quantified as described 275 previously (Rivas et al. 2011; Blanco et al. 2012). An additional cut-off threshold, which was the 276 interval of F with an accumulated frequency higher than 80 %, was applied to minimize the

277	probability of false positives. Results were finally represented as a color-patterned heatmap for a
278	better visualization, normalizing output $F$ attending to number of antibodies per
279	taxonomic/metabolic group and to total microarray fluorescence values (He et al. 2007).
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281	3. Results
282	
283	3.1. MinION sequencing produced diverse metagenomes from the Hanksville
284	paleochannels
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286	DNA was extracted from all samples and sequenced with the MinION (Table 1). The
287	highest input was 19.5 ng from sample R2 and the lowest was 0.667 ng from sample S2. With
288	the exceptions of samples E1, S3, and R2, total reads produced generally increased with input
289	DNA, while average sequence length was between 2598 bp and 3571 bp for all samples. After
290	adaptor trimming and correction with Porechop and Canu, the total number of reads decreased
291	sharply as a result of Canu's error correction, removal of noisy sequences, and replacement with
292	consensus sequences. The number of hits classified by JGI IMG/M ER was very low (0.72 $\%-$
293	9.15 %) (Table 1). This is not uncommon for nanopore sequencing, which can have classification
294	rates ranging from 15 % (Edwards et al. 2017) single strand 1D sequencing and up to 98 %
295	(Brown <i>et al.</i> 2017) for double strand for $1D^2$ sequencing. This may be partially due to high error
296	rates in MinION basecalling (Jain et al. 2017) and is likely exacerbated in natural microbial
297	consortia. Additionally, metagenomic datasets tend to have low taxon classification in general
298	due to incomplete databases and higher numbers of unclassifiable sequences, as opposed to
299	amplicon sequencing datasets (Tessler et al. 2017).

300 The Hanksville paleochannel metagenomes contain rich microbial communities, with 301 members from all three domains of life (Bacteria, Archaea, and Eukarya), as well as viruses (Fig. 302 2). The samples contain a wide variety of Bacteria, with samples E1 and R1 each containing 303 genes from 21 distinct bacterial phyla. Proteobacteria, Actinobacteria, Cyanobacteria, and 304 Bacteroidetes were most prominent phyla encountered. Proteobacteria and Actinobacteria were 305 detected in all 11 samples, and in large proportions ( $\geq 10$  %) in 7 and 3 samples, respectively. 306 Three endolith samples (E3, E4, and E5) were strongly dominated by Cyanobacteria ( $\geq$  76 %), 307 followed by Bacteroidetes and Proteobacteria. Nostocales was the most prominent 308 cyanobacterial order present. Among the sediment and rock samples, Actinobacteria, 309 Bacteroidetes, and Proteobacteria were the dominant taxa, save sample R3, which had 310 Cyanobacteria as a chief phylum (38%). This sample had also had Nostocales as its most 311 abundant cyanobacterial order. R3 had no visible endoliths and instead was covered by a dark oxidized coating (likely desert varnish). Desert varnish has been observed at the Viking and 312 313 Pathfinder landing sites on Mars (McCauley et al. 1979; Guinness et al. 1997), and while the 314 method of desert varnish formation remains unclear, it is postulated to involve microbial 315 mediation and deposition of manganese-rich films (Dorn and Oberlander 1981). Archaea were 316 present in small proportions (~1%) in 7 out of 11 samples as Euryarchaeota and 317 Thaumarchaeota. Among the bacterial and archaeal genes detected are several extremophiles 318 capable of surviving Mars-like conditions, further enforcing the analogue potential of this site 319 beyond geomorphology. These extremophiles include UV-resistant and halophilic Halobacteria, 320 methanogens Methanobacteria and Methanomicrobia, desiccation-tolerant Chroococcidiopsis, 321 and 25 distinct genomes of the polyextremophilic phylum Deinococcus-Thermus (Direito et al. 322 2011). Also present in the paleochannel metagenomes is a diverse range of eukaryotes, including

fungi as Ascomycota, algae as Streptophyta and Chlorophyta, and unclassified Eukarya (mainly
algal protists and choanoflagellates).

325

326 3.2. Illumina 16S rRNA gene sequencing revealed a rich bacterial community in the
 327 Hanksville paleochannels

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329 Given the large proportion of bacteria detected in the MinION sequencing data, we 330 performed 16S rRNA gene sequencing to better illuminate the bacterial diversity present and 331 provide a comparison to the MinION results. This comparison can be seen in Fig. 3A and Fig. 332 3B. Sequencing the 16S rRNA gene content of the 11 samples yielded 2 478 778 reads of mean 333 length 253 bp. 1435 OTUs were generated after clustering at 97 % similarity. The most abundant 334 phyla detected among these OTUs were Actinobacteria, Cyanobacteria, Proteobacteria, 335 Deinococcus-Thermus, and Chloroflexi (Fig. 3A). and these were detected in all samples. In 336 particular, Actinobacteria comprised > 5 % of OTUs in all samples. Like the MinION results, 337 three of the endolith samples were strongly dominated by Cyanobacteria (particularly 338 Nostocales) ( $\geq$  44 %) but were samples E1, E2, and E4 (rather than E3, E4, and E5). Unlike the 339 MinION data, E3 and E5 contained higher proportions of Actinobacteria and Deinococcus-340 Thermus (mostly Deinococcales) ( $\geq$  13 %) instead of Cyanobacteria. Sediment and rock samples 341 were dominated by Actinobacteria, Proteobacteria, and Chloroflexi. The endolith, sediment, and 342 rock samples showed some differences in taxonomy, which was visualized with a weighted 343 UniFrac principal coordinates analysis (PCoA) of the 16S data (PERMANOVA: pseudo-f 344 statistic = 2.89990; p-value = 0.01905) (Anderson and Walsh 2013) (Fig. 4). The endoliths

appear to cluster solely with other endolith samples (E1, E2, and E4 vs. E3 and E5),

- 346 demonstrating the closer phylogenetic relationship within the endolith samples.
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348 *3.3. Determination of putative DNA detection limits for MinION sequencing* 

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350 We used high quality DNA from endolith sample E4 to determine prospective DNA 351 detection limits of MinION sequencing (Table 2; Fig. 5). We began our detection limit trials at 352 0.110 ng of DNA, a tenth of the minimum recommended input from ONT. Sequencing 0.110 ng 353 of DNA from E4 yielded 253 188 total reads, nearly 26 times more than when 4.04 ng of DNA 354 from this sample was sequenced. This is likely due to the use of R9.4 FLO-MIN106 flow cells 355 rather than R9.5 FLO-MIN107 flow cells. R9.5 chemistry was initially recommended during our 356 metagenome sequencing but has since been advised for only  $1D^2$  double strand sequencing by 357 ONT (1D single strand sequencing was employed in this study). We subsequently tested 0.0408 358 ng, 0.0106 ng, and 0.00106 ng of DNA. Total read number dropped with input concentration, as 359 expected, and the metagenomes produced during these trials are relatively similar (Fig. 5). 360 Cyanobacteria are the dominant phylum, followed by Proteobacteria and Bacteroidetes. Other 361 phyla become more prominent as input DNA decreases and Cyanobacteria are completely 362 undetected in the final 0.00106 ng trial. We also determined if the MinION would produce a false positive with no DNA present. Sequencing nuclease-free water containing no DNA 363 364 produced 121 reads, likely as a result of barcode adapter dimers and PCR amplification of these 365 dimers. None of these reads were classified by JGI IMG/M ER. We also examined the quality 366 score distribution of this run, as compared to the other DNA inputs in the detection limit trials 367 (Supplementary Fig. 1). PHRED score is a common metric of sequence quality (Manley et al.

368 2016) and decreases strongly with DNA input in our study, with a seemingly random distribution 369 in our run containing no DNA. Low sequence number, lack of classification, and random quality 370 score distributions are therefore strongly indicative of a run in which no environmental DNA is 371 present. 372 373 3.4. Powder X-ray Diffraction (XRD) revealed a quartz-rich environment 374 375 The X-ray diffraction patterns of the Hanksville paleochannel samples are dominated by 376 quartz (Supplementary Table 1). Many samples also contain minor chemical sedimentary 377 minerals (gypsum and calcite) and a phyllosilicate component identified in most cases as 378 montmorillonite. The diffraction patterns for endolith samples E2, E3, and E4 contain only 379 quartz, while endoliths E1 and E5 also exhibit traces of calcite and albite (possible Ca-rich), 380 respectively. Regolith samples S1 and S2 display similar mineralogy, with minor of albite (Ca-381 rich) and montmorillonite. Regolith S3 is more heterogenous and contains calcite, gypsum, albite 382 (Ca-rich), and montmorillonite. The non-endolithic rock samples R1, R2, and R3 are diverse; R2 383 is composed of minor of magnesian calcite, carbonate-fluorapatite, and nontronite. R1 contains 384 montmorillonite, gypsum, and apatite-(CaCl), while R3 has a minor calcite component. 385 386 3.5. Reflectance spectra show hydrated minerals and some organic features 387 388 The reflectance spectra of all samples are dominated by absorption features characteristic 389 of hydrated mineral phases (Fig. 6). Prominent absorption features include overtones and 390 combinations of H<sub>2</sub>O and OH<sup>-</sup> stretches and bends occurring at ~ 1400 nm (OH<sup>-</sup>), ~ 1900 nm

391	(H <sub>2</sub> O and OH <sup>-</sup> ), and a combination of vibrational stretching and bending of Al <sub>2</sub> OH at ~ 2220 nm.
392	The shape and position of the main ~ 1400, ~ 1900, ~ 2200 nm (Al-OH), and 2350 nm (Mg-OH)
393	features in the samples provide insights into the hydrated and sheet silicate phases present. The
394	presence of a ~ 1400 nm and ~ 1900 nm band indicates that both $OH^-$ and molecular water are
395	present in the samples (Cloutis et al. 2006). The position, shape and depth of these water features
396	and the ~ 2200 nm (band center at ~ 2212 nm) feature indicates that the bulk of the samples have
397	spectra characteristic of montmorillonite (Clark et al. 1990; Bishop et al. 2008). The weak
398	feature observed in multiple samples at ~ 2300 nm is a C-O feature due to the second overtone of
399	a CO <sub>2</sub> asymmetrical stretch and is indicative of carbonate minerals (Fig. 6B) (Clark et al. 1990).
400	This position is indicative of carbonate chemistry and in these samples supports the presence of
401	calcite. Absorption features of Fe-bearing oxyhydroxyides such as goethite at ~ 500 nm, ~ 900
402	nm, and ~ 1150 nm are also evident, as is the characteristic drop in reflectance below ~ 1000 nm
403	due to charge transfers in Fe-bearing minerals (Crowley et al. 2003).
404	Organic features are also present, primarily in samples containing visible endoliths (E1,
405	E2, E3, E4, and E5). These samples exhibit a characteristic absorption band at 670 nm, a sharp
406	increase in reflectance ("chlorophyll bump" or "green peak") preceding this feature, and an
407	overall drop in reflectance below ~ 800 nm ("red edge") that are all diagnostic of chlorophyll a
408	(Rhind et al. 2014; Stromberg et al. 2014). A small peak at ~ 400 nm can be attributed to the
409	presence of porphyrins, resulting either from the degradation of chlorophylls or as a component
410	of various cytochromes (Berg et al. 2014). A broad band at ~ 465 nm can be indicative of either
411	carotenoids (Berg et al. 2014) or NADPH (Ammor 2007). A generic C-H overtone at ~ 1750 nm
412	is also present, but N-bearing molecules can also affect absorption features in this region (Berg et
413	<i>al.</i> 2014).

414

- 415 *3.6. Prominent bands of chlorophyll and quartz are observed in the Raman spectra*
- 416

417	The samples exhibit a large number of distinct Raman peaks at an excitation wavelength
418	of 532 nm (Fig. 7), which was chosen to optimize biomolecule detection and mimic that of the
419	Raman Laser Spectrometer (RLS) on ExoMars 2020 (Edwards et al. 2012; Rull et al. 2017).
420	Prominent peaks associated with the mineralogy of the site are observed at ~ $470 \text{ cm}^{-1}$ (quartz)
421	and ~ 1006 cm <sup>-1</sup> (gypsum) (Edwards <i>et al.</i> 2007). The non-endolith samples are dominated by
422	background iron fluorescence, a common artifact in Fe-containing samples and a 532 nm source.
423	Other noticeable features include a small peak at ~ $930 \text{ cm}^{-1}$ , which is attributable to C-COO <sup>-</sup>
424	stretching vibrations (Kim et al. 1987) or the presence of perchlorate (Jones et al. 1961; Ruan et
425	al. 2006). Asymmetric stretching vibrations from $CO_2$ are presented as a small band at ~ 2340
426	cm <sup>-1</sup> (Sandford and Allamandola 1990; Feng <i>et al.</i> 2004). A peak at ~ 1655 cm <sup>-1</sup> is indicative of
427	the presence of organics and represents the amide I band in classical Raman spectroscopy. This
428	band is the result of asymmetric C=O stretching vibrations and $\alpha$ -helical secondary structure in
429	proteins (Thomas et al. 1987; Maiti et al. 2004).
430	The endolith-bearing samples also contain numerous features indicative of the presence

<sup>430</sup> The endoluth-bearing samples also contain numerous features indicative of the presence <sup>431</sup> of organic matter. A broad spectral feature at ~ 3530 cm<sup>-1</sup> signifies the presence of chlorophyll in <sup>432</sup> these samples (Rhind *et al.* 2014). Peaks at ~ 1155 cm<sup>-1</sup> and ~ 1520 cm<sup>-1</sup> are the result of C-C <sup>433</sup> and C=C stretching vibrations, respectively, in various carotenoids, including β-carotene (Ellery <sup>434</sup> *et al.* 2002; Edwards *et al.* 2005; Rhind *et al.* 2014), bacteriorhodopsin (Jehlička *et al.* 2013), <sup>435</sup> bacterioruberin (Marshall *et al.* 2006), astaxanthin (Kaczor and Baranska 2011), and lutein <sup>436</sup> (Edwards *et al.* 2007). Additionally, while the ~ 470 cm<sup>-1</sup> band is characteristic of quartz, this

437	feature can also indicate the presence of parietin, a secondary UV-protectant pigment. Edwards
438	et al. (2005) noted that the ~ 470 cm <sup>-1</sup> feature only appeared in samples from known parietin-
439	containing environments and attributed its cause to parietin rather than quartz.
440	
441	3.7. LDChip immunoassays detected a wide variety of potential metabolisms and taxa
442	
443	Positive signals corresponding to antibodies against numerous bacterial, archaeal, and
444	viral organisms were detected by the LDChip (Fig. 8 and 9). The strongest immunoassay signals
445	corresponded to biomarkers associated to Actinobacteria (Streptomyces), Cyanobacteria
446	(Aphanizomenon, Chroococcidiopsis), iron-sulfur bacteria from Acidithiobacillia
447	(Acidithiobacillus), Nitrospirae (Leptospirillum), and crude environmental extracts from the Río
448	Tinto that are rich in iron and sulfur. Proteins relating to common microbial functions (e.g.
449	chaperones HscA2 and HtpG) were readily detected, as well as more specialized functions, such
450	as nitrogen metabolism (e.g. NifD and nitrate reductase) and hydric stress (e.g. DhnA). The
451	LDChip also detected proteins related to perchlorate reduction in samples E1, E2, S1, and S3.
452	Archaeal markers associated with halophiles (Haloferax) and methanogens (Methanobacterium)
453	were detected in samples E3, and E4, S1, S2, S3, and R1, respectively.
454	
455	4. Discussion
456	
457	4.1. MinION metagenomes complement and support XRD data, reflectance spectra,
458	Raman spectra, and LDChip immunoassays
459	

460 Here we have shown that the MinION is capable of detecting a wide variety of 461 microorganisms from all three domains of life, as well as viruses (Fig. 2). Our results are 462 consistent with a previous study by Direito et al. (2011), who conducted a 16S rRNA survey of 463 the Mars Desert Research Station (MDRS) soil, an area adjacent to the Hanksville paleochannels. Direito et al. (2011) found overall high diversity in the bacterial and eukaryal 464 465 populations of this area, and low archaeal diversity. Bacteria dominated the total microbial content in our study, as well as in a PCR-based detection survey of the MDRS (Thiel et al. 466 467 2011). Strangely, we also detected *Mollusca*, *Cnidaria*, and *Porifera*. These organisms are not 468 common laboratory contaminants, but Mollusca were present in the Direito et al. (2011) study 469 where MinION sequencing was not used. Direito et al. (2011) noted the presence of a hill 470 consisting fossil oyster shells in this area, signifying the marine history of this environment and 471 indicating that these sequences are potentially naturally occurring. 472 Direct comparison of the bacterial taxa detected in both Illumina and MinION sequencing 473 shows some overlap (Fig. 3A and Fig. 3B, respectively). In particular, Acidobacteria, 474 Actinobacteria, Bacteroidetes, Chloroflexi, Cyanobacteria, Planctomycetes, and Proteobacteria 475 were detected in proportions > 15 % for all reads in both runs. While shotgun metagenome 476 datasets often reveal more diversity and more rare taxa in an environment (Poretsky et al. 2014; 477 Ranjan et al. 2016), our 16S amplicons revealed more diversity in the Hanksville paleochannels 478 than our shotgun MinION data. Specifically, the candidate phylum FBP was detected in 7 of our 479 11 samples in the amplicon dataset but was completely absent in our MinION data. FBP is 480 closely related to Armatimonadetes (Lee et al. 2013) and is a common member of extreme 481 environmental consortia (Tahon et al. 2018). Gemmatimonadetes was also a significant 482 component of our 16S dataset (mainly Longimicrobiales) but was detected only as small

483	amounts of Gemmatimonadales in E3 and E5 in our shotgun dataset. This discrepancy was also
484	observed by Tessler <i>et al.</i> (2017), who found that $< 50$ % of phyla detected in 16S amplicon
485	datasets were also found in shotgun datasets. Conversely, nearly all phyla detected in their
486	shotgun metagenomes were also picked up in their amplicon data. Inherent differences between
487	these techniques can account for these differences, namely reduced database size for shotgun
488	metagenomes and a higher proportion of reads mapped to unknown taxa in shotgun datasets
489	(Tessler et al. 2017). Our study introduces additional challenges that can contribute to the lower
490	diversity observed in our shotgun data: low read output and the MinION's inherently high error
491	rate of ~ 7.5 % – 14.5 % (Jain <i>et al.</i> 2017). As a comparison, Illumina sequencing has an error
492	rate of < 0.1 % (Manley <i>et al.</i> 2016).
493	The LDChip and MinION revealed somewhat different microbial profiles, both

494 phylogenetically and metabolically (Fig. 8 and 9). While both sets of data detected 495 Cyanobacteria, Proteobacteria, Bacteroidetes, and Firmicutes, differences arose at subphyla 496 levels. The intrinsic differences in each technique can account for these variations. The groups 497 detected by the LDChip can be very broad, while MinION sequencing is inherently specific to a 498 certain gene or taxa. While the LDChip's polyclonal antibodies are raised to detect certain 499 proteins and peptides, they might be also recognizing similar epitopes in other proteins or closely 500 related taxa (Parro et al. 2011). Additionally, the specific microorganism may have been present 501 but not had the gene in question sequenced by the MinION. The LDChip can also detect dead 502 microbial polymers (e.g. proteins and EPS) even when neither cells nor DNA are present. This 503 indicates the powerful complementarity of both techniques to detect a wide variety of 504 microorganisms and different types of definitive biosignatures in an environment (proteins and 505 nucleic acids).

506 XRD results (Supplementary Table 1), reflectance spectra (Fig. 6), and Raman spectra 507 (Fig. 7) provide geological context for our MinION data and characterize the surrounding 508 environment. LDChip results (Fig. 8 and 9) also supplement the MinION sequences through 509 more targeted detection of specific molecular markers. Table 4 shows the complementary 510 relationship between these methods. The Hanksville paleochannel environment is quartz-rich, as 511 all samples presented quartz as their major mineral detected by XRD. Quartz has been shown to 512 support cyanobacterial desert communities in previous studies (Schlesinger et al. 2003; Warren-Rhodes et al. 2006; Lacap et al. 2011). This corroborates the dominance of Cyanobacteria in 513 514 samples E3, E4, E5, and R3 in the MinION metagenomes, as well as the presence of 515 Cyanobacteria in all samples in both the MinION and LDChip data. Raman spectra also revealed 516 quartz as a ~  $470 \text{ cm}^{-1}$  peak. Calcite was present as a minor mineral in the XRD spectra, and is 517 known to be commonly precipitated by many types of environmental bacteria (Boquet et al. 1973; Castanier et al. 1999), including sulfate-reducing bacteria (SRBs) (Van Lith et al. 2003) 518 519 and various species of Bacillus (Boquet et al. 1973). Bacillus were detected in our MinION 520 sequences, as were SRBs from Desulfobacterales and Desulfovibrionales. The presence of SRBs 521 was further indicated by genes involved in sulfate transport, sulfate-sulfur assimilation, and 522 sulfur and sulfate metabolism (e.g. sulfate transport system permease and sulfate 523 adenylyltransferase). The LDChip also detected Bacillus and SRBs among the Firmicutes, 524 Deltaproteobacteria, and positive results for antibodies raised against crude environmental 525 extracts from the Río Tinto that are rich in iron and sulfur. Gypsum was in the samples, as indicated by XRD spectra and a Raman peak at ~ 1006  $\text{cm}^{-1}$  and gypsum itself a source of 526 527 sulfate. It therefore could host the same SRBs and their associated protein groups. Gypsum is 528 also conducive to the presence of archaeal and cyanobacterial halophiles (Oren et al. 2009;

530 C	Cyanobacteria and Halobacteria by the MinION and LDChip. Iron was detected as Fe-bearing
531 o	oxyhydroxides in the reflectance spectra and as background fluorescence in the Raman spectra.
532 I	fron is crucial for most bacterial metabolisms (Frawley and Fang 2014), and Fe-oxyhydroxides
533 a	are associated with both iron-oxidizing bacteria (Emerson et al. 1999) and iron-reducing bacteria
534 (	(Luef et al. 2013). The LDChip exhibited positive results for antibodies raised against crude
535 e	environmental extracts rich in iron and sulfur from the Río Tinto in all samples. Heme oxygenase
536 v	was detected in our MinION data, as well. Perchlorate may have been responsible for the ~ 930
537 c	cm <sup>-1</sup> feature in non-endolith Raman spectra and perchlorate reductase was detected by the
538 L	LDChip. Perchlorate comprises $0.5 - 1$ % of mid-latitude Martian surface soils (Hecht <i>et al.</i>
539 2	2009; Glavin et al. 2013), and perchlorate reduction genes indicate the presence of
540 n	nicroorganisms that can tolerate perchlorate-rich soils and that could serve as models for
541 N	Martian life. Perchlorate can also be used as a terminal electron acceptor during anaerobic
542 g	growth for perchlorate-reducing bacteria (Coates et al. 1999).
543	We also detected numerous signatures for biological pigments and other metabolisms. In
544 tl	he reflectance spectra, chlorophyll is presented as a characteristic absorption band at 670 nm, its
545 p	preceding "chlorophyll bump", and its trailing "red edge" (Rhind et al. 2014; Stromberg et al.
546 2	2014). Porphyrins are denoted by a small peak at ~ 400 nm (Giovannetti 2012; Berg et al. 2014),
547 a	and a broad band at ~ 465 nm can be indicative of carotenoids (Berg et al. 2014). Raman spectra
548 a	also contain a strong chlorophyll feature at ~ 3530 cm <sup>-1</sup> , and peaks at ~ 1155 cm <sup>-1</sup> and ~ 1520
549 c	cm <sup>-1</sup> indicate the presence of various carotenoids, including $\beta$ -carotene (Ellery <i>et al.</i> 2002;
550 E	Edwards et al. 2005; Rhind et al. 2014), bacteriorhodopsin (Jehlička et al. 2013), bacterioruberin
551 (	Marshall <i>et al.</i> 2006), astaxanthin (Kaczor and Baranska 2011), and lutein (Edwards <i>et al.</i>

552 2007). Parietin may also be denoted by a band at ~ 470 cm<sup>-1</sup> (Edwards et al. 2005). The presence 553 of microbial pigments is supported by both MinION and LDChip results, which detected 554 Cyanobacteria in all samples. Specifically, genes encoding chlorophyll's photosystem I and II, 555 protochlorophyllide reductase,  $\beta$ -carotene ketolase and hydroxylase, and protoporphyrin IX 556 cyclase were detected by the MinION. The LDChip indicated the presence of nitrogen 557 metabolizing microorganisms, as did some MinION sequences, but they differed on the specific 558 gene or protein. The LDChip detected NifD1 and nitrate reductase, while the MinION detected 559 NifU and NirS. Together, the LDChip and the MinION were able to detect a wider variety of 560 genes involved in the nitrogen cycle than either alone. The same is true for methanogenic 561 microorganisms. *Methanobacterium* was detected by the LDChip, while *Methanosarcina*, 562 Methanothrix, Methanolacinia, and Methanosarcinales were detected by the MinION. Through 563 this combination of multiple techniques, we have characterized this paleochannel environment as 564 rich in quartz, calcite, and gypsum and able to host a wide variety of microorganisms, including 565 numerous Cyanobacteria and SRBs. We can also infer these microbes' potential metabolisms, 566 which includes nitrogen cycling, sulfur cycling, methane production, and perchlorate reduction. 567 568 4.2. Putative MinION DNA detection limits and suitability for life detection 569

570 DNA is an ideal target in the search for life since it is unambiguous, non-specific, and 571 readily detectable with nanopore sequencing. Past meteoritic exchange between Earth and Mars 572 (Mileikowsky *et al.* 2000), as well as the abundance of carbon molecules in our Solar System 573 (Pace 2001) indicates a plausibility of organic-based biochemistry for potential Martian life 574 (Mustard *et al.* 2013b). However, an extremely important criterion for life detection in sample

575 return and other space exploration missions is the ability to detect low biomass, such as that 576 present in various Mars analogue environments. We have successfully shown that the MinION 577 can detect DNA as low as  $\sim 0.001$  ng and that it will not produce a false positive when no DNA 578 is present. This is further enforced by Pontefract et al. (2018), who showed that the MinION will 579 not produce sequencing artifacts or noise signals that could be mistaken positive results. 580 Producing sequencing data in real-time is also a strong asset. Chiefly, it would enable the best 581 selection of samples to cache for potential future return to Earth (e.g. those containing DNA or 582 other nucleic acids) (Goordial *et al.* 2017). While the MinION's high error rate ( $\sim 7.5 \% - 14.5$ 583 %) (Jain et al. 2017) and low classification rate (Edwards et al. 2016) can be problematic, a well-584 curated genetic database of clean room contaminants would enable the quick and efficient 585 identification of contamination on missions. Our comparison between Illumina 16S amplicon 586 data and MinION shotgun metagenome data has shown that the MinION is able to reliably detect 587 common lineages in an environment. Despite the possibility of shared heritage between 588 terrestrial and potential Martian organisms (Mileikowsky et al. 2000; Carr et al. 2017), the 589 divergence between them would likely have occurred ~ 4 billion years ago and any Martian 590 DNA would map outside of modern branches on the tree of life (Fairén et al. 2017). Therefore, 591 any potential Martian life would be an independent lineage and unable to be classified. The 592 MinION has been shown to sequence inosine in nucleic acid chains (Carr *et al.* 2017), 593 demonstrating its ability to detect non-standard DNA and indicating its potential to detect a 594 wider variety of charged chains. This indicates its potential for life detection on non-Earth like 595 worlds, such as the gas giants and icy moons (e.g. Enceladus and Europa) (Mustard et al. 2013b). 596 In fact, a solid state (i.e. non-degradable) nanopore sequencer for detection of organic linear

polymers is currently being developed by NASA's Concepts for Ocean worlds Life Detection
Technology program (Bywaters *et al.* 2017).

599 While low input preparations for MinION sequencing (such as that employed here) 600 require a PCR step for amplification and adapter attachment, the primers used are non-specific, 601 and instead are attached to nucleic acid strands via transposase mediation. This greatly reduces 602 the inherent bias present in PCR and renders the overall protocol non-specific for certain genes 603 or organisms. However, this introduces the need for a thermocycler in MinION-based life 604 detection. PCR is possible in space and is able to be miniaturized for deployment, as shown by 605 the successful use of the miniPCR system aboard the ISS during the Biomolecule Sequencing 606 project (John *et al.* 2016), but it is one more instrument that needs to be automated and added to 607 the sequencing process.

608 Our putative lower limit for sequencing is  $\sim 0.001$  ng of DNA, which corresponds to 100 609 cells/g (Dolezel et al. 2003; Raes et al. 2007), a biomass common in Mars analogue 610 environments (Niederberger et al. 2010; Goordial et al. 2016). This is based on an average 611 bacterial genome size of 2-5 Mbp and  $\sim 2.5-5$  fg of DNA per genome (Dolezal *et al.* 2003; 612 Raes et al. 2007; Goordial et al. 2017). However, this assumes 100 % extraction efficiency, 613 which is unlikely. The lower limit of DNA detection in a mission will be strongly dependent on 614 the extraction method and must be chosen to optimize efficiency, as well as automation and 615 miniaturization. Sequencing preparation will also need to be automated and miniaturized before 616 deployment. This could be accomplished with the SOLID-SPU (Signs of Life Detector Sample 617 Preparation Unit) (Parro et al. 2011) for DNA extraction and the VolTRAX device by ONT for 618 sequencing preparation. Both instruments require relatively little human intervention but would

still require a microfluidics system to connect the components. We are exploring these optionspresently.

621 The techniques we have employed here form a powerful suite capable of biosignature 622 detection and environmental characterization. The MinION was able to directly detect 623 unambiguous biosignatures in all sample types (endolith, regolith, rock) from this paleochannel site analogous to Martian sinuous ridges. There are prominent organic and water-related features 624 625 in the XRD data, reflectance spectra, and Raman spectra, though by themselves these 626 biosignatures are not conclusive and an abiotic origin cannot be precluded. Features indicating 627 pigments are present but only in samples containing high biomass endoliths. The LDChip 628 detected signals corresponding only to the antibodies present on the chip and this introduces a 629 strong bias towards very specific protein groups and species. However, when these techniques 630 are combined with the MinION, which can detect and sequence DNA from a wide variety of 631 microorganisms in many types of samples and in very small amounts, the potential for 632 biosignature identification and characterization is greatly increased. 633 634 **5.** Conclusion 635 636 We present this work as a proof-of-concept that MinION sequencing is a promising 637 candidate for direct life detection in future space exploration missions. It can detect DNA down 638 to  $\sim 0.001$  ng (and potentially lower). The MinION is small, portable, low cost, and can

639 successfully detect life in many analogue environments, including the Hanksville paleochannels.

640 Dominated by Bacteria, the taxa detected here were able to be linked to the surrounding

641 environment with XRD data, reflectance spectroscopy, Raman spectroscopy, and LDChip

642	immunoassays. We can infer the presence of Cyanobacteria and SRBs based on these
643	complementary analyses. Nanopore technology is synergistic with these techniques and together
644	they provide a thorough characterization of this environment and its biosignatures. Although
645	development of the MinION as a standalone instrument offers significant challenges (e.g.
646	automated nucleic acid extraction and sample preparation), it represents a positive future for a
647	direct life detection payload.
648	
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650	
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657	
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660	No competing financial interests exist.
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665 **References** 

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668	identification and characterization. Journal of fluorescence, 17: 455-459.
669	Anderson M. J., and Walsh D. C. (2013) PERMANOVA, ANOSIM, and the Mantel test in the
670	face of heterogeneous dispersions: what null hypothesis are you testing? Ecological
671	monographs, 83: 557-574.
672	Balme M., Grindrod P., Sefton-Nash E., Davis J., Gupta S., Fawdon P., Sidiropoulos P., Yershov
673	V., and Muller JP. (2015) Aram Dorsum: A Noachian Inverted Fluvial Channel System
674	and Candidate Exomars 2018 Rover Landing Site.Lunar and Planetary Science

Ammor M. S. (2007) Recent advances in the use of intrinsic fluorescence for bacterial

675 Conference.

676 Beegle L., Bhartia R., White M., DeFlores L., Abbey W., Wu Y.-H., Cameron B., Moore J.,

677 Fries M., and Burton A. (2015) SHERLOC: Scanning habitable environments with

- Raman & luminescence for organics & chemicals.Aerospace Conference, 2015 IEEE.
  IEEE.
- 680 Berg B., Ronholm J., Applin D., Mann P., Izawa M., Cloutis E., and Whyte L. (2014) Spectral

features of biogenic calcium carbonates and implications for astrobiology. *International Journal of Astrobiology*, 13: 353-365.

Bishop J. L., Dobrea E. Z. N., McKeown N. K., Parente M., Ehlmann B. L., Michalski J. R.,

- Milliken R. E., Poulet F., Swayze G. A., and Mustard J. F. (2008) Phyllosilicate diversity
  and past aqueous activity revealed at Mawrth Vallis, Mars. *Science*, 321: 830-833.
- 686 Blanco Y., Prieto-Ballesteros O., Gómez M. J., Moreno-Paz M., García-Villadangos M.,
- 687 Rodríguez-Manfredi J. A., Cruz-Gil P., Sánchez-Román M., Rivas L. A., and Parro V.

- (2012) Prokaryotic communities and operating metabolisms in the surface and the
  permafrost of Deception Island (Antarctica). *Environmental microbiology*, 14: 24952510.
- Boquet E., Boronat A., and Ramos-Cormenzana A. (1973) Production of calcite (calcium
  carbonate) crystals by soil bacteria is a general phenomenon. *Nature*, 246: 527.
- Brown B. L., Watson M., Minot S. S., Rivera M. C., and Franklin R. B. (2017) MinION<sup>™</sup>
- 694 nanopore sequencing of environmental metagenomes: a synthetic approach. *GigaScience*,695 6: 1-10.
- 696 Bywaters K., Schmidt H., Vercoutere W., Deamer D., Hawkins A., Quinn R., Burton A., and

697 Mckay C. (2017) Development of Solid-State Nanopore Technology for Life Detection.

698 Carr C. E., Mojarro A., Hachey J., Saboda K., Tani J., Bhattaru S. A., Smith A., Pontefract A.,

EVALUATE: Comparison of the Second Se

700 detection.Aerospace Conference, 2017 IEEE. IEEE.

701 Castanier S., Le Métayer-Levrel G., and Perthuisot J.-P. (1999) Ca-carbonates precipitation and

702 limestone genesis—the microbiogeologist point of view. *Sedimentary geology*, 126: 9-23.

703 Castro-Wallace S. L., Chiu C. Y., John K. K., Stahl S. E., Rubins K. H., McIntyre A. B.,

704 Dworkin J. P., Lupisella M. L., Smith D. J., and Botkin D. J. (2017) Nanopore DNA

sequencing and genome assembly on the International Space Station. *Scientific reports*,

- 706 7: 18022.
- Clark R. N., King T. V., Klejwa M., Swayze G. A., and Vergo N. (1990) High spectral resolution
   reflectance spectroscopy of minerals. *Journal of Geophysical Research: Solid Earth*, 95:
   12653-12680.

- 710 Clarke J. D., and Stoker C. R. (2011) Concretions in exhumed and inverted channels near
- Hanksville Utah: implications for Mars. *International Journal of Astrobiology*, 10: 161175.
- 713 Cloutis E. A., Hawthorne F. C., Mertzman S. A., Krenn K., Craig M. A., Marcino D., Methot M.,
- Strong J., Mustard J. F., and Blaney D. L. (2006) Detection and discrimination of sulfate
  minerals using reflectance spectroscopy. *Icarus*, 184: 121-157.
- 716 Coates J. D., Michaelidou U., Bruce R. A., O'Connor S. M., Crespi J. N., and Achenbach L. A.
- 717 (1999) Ubiquity and diversity of dissimilatory (per) chlorate-reducing bacteria. *Applied*
- 718 *and Environmental Microbiology*, 65: 5234-5241.
- 719 Crowley J., Williams D., Hammarstrom J., Piatak N., Chou I.-M., and Mars J. (2003) Spectral
- 720 reflectance properties (0.4–2.5 μm) of secondary Fe-oxide, Fe-hydroxide, and Fe-
- sulphate-hydrate minerals associated with sulphide-bearing mine wastes. *Geochemistry: Exploration, Environment, Analysis*, 3: 219-228.
- Davis J., Balme M., Grindrod P., Williams R., and Gupta S. (2016) Extensive Noachian fluvial
   systems in Arabia Terra: Implications for early Martian climate. *Geology*, 44: 847-850.
- 725 Direito S. O., Ehrenfreund P., Marees A., Staats M., Foing B., and Röling W. F. (2011) A wide
- variety of putative extremophiles and large beta-diversity at the Mars Desert Research
  Station (Utah). *International Journal of Astrobiology*, 10: 191-207.
- Dolezel J., Bartos J., Voglmayr H., and Greilhuber J. (2003) Nuclear DNA content and genome
  size of trout and human. *Cytometry. Part A: the journal of the International Society for Analytical Cytology*, 51: 127.
- Dorn R. I., and Oberlander T. M. (1981) Microbial origin of desert varnish. *Science*, 213: 12451247.

- 733 Downs R. (2006) The RRUFF Project: an integrated study of the chemistry, crystallography,
- Raman and infrared spectroscopy of minerals.Program and Abstracts of the 19th General
- 735 Meeting of the International Mineralogical Association in Kobe, Japan, 2006.
- 736 Edwards A., Debbonaire A. R., Sattler B., Mur L. A., and Hodson A. J. (2016) Extreme
- 737 metagenomics using nanopore DNA sequencing: a field report from Svalbard, 78 N.
- *bioRxiv*: 073965.
- 739 Edwards A., Soares A., Rassner S., Green P., Felix J., and Mitchell A. (2017) Deep Sequencing:

740 Intra-terrestrial metagenomics illustrates the potential of off-grid Nanopore DNA

741 sequencing. *bioRxiv*: 133413.

- Edwards H. G., Hutchinson I., and Ingley R. (2012) The ExoMars Raman spectrometer and the
- 743 identification of biogeological spectroscopic signatures using a flight-like prototype.

744 *Analytical and bioanalytical chemistry*, 404: 1723-1731.

- 745 Edwards H. G., Vandenabeele P., Jorge-Villar S. E., Carter E. A., Perez F. R., and Hargreaves
- 746 M. D. (2007) The Rio Tinto Mars analogue site: an extremophilic Raman spectroscopic
- 747study. Spectrochimica Acta Part A: Molecular and Biomolecular Spectroscopy, 68: 1133-
- 748 1137.
- Edwards H. G., Villar S. E. J., Parnell J., Cockell C. S., and Lee P. (2005) Raman spectroscopic
  analysis of cyanobacterial gypsum halotrophs and relevance for sulfate deposits on Mars. *Analyst*, 130: 917-923.
- Ehrenfreund P., Röling W., Thiel C., Quinn R., Sephton M., Stoker C., Kotler J., Direito S.,
- 753 Martins Z., and Orzechowska G. (2011) Astrobiology and habitability studies in
- 754 preparation for future Mars missions: trends from investigating minerals, organics and
- biota. International Journal of Astrobiology, 10: 239-253.

- 756 Eigenbrode J. L., Summons R. E., Steele A., Freissinet C., Millan M., Navarro-González R.,
- 757 Sutter B., McAdam A. C., Franz H. B., and Glavin D. P. (2018) Organic matter preserved
- in 3-billion-year-old mudstones at Gale crater, Mars. *Science*, 360: 1096-1101.
- Ellery A., Kolb C., Lammer H., Parnell J., Edwards H., Richter L., Patel M., Romstedt J.,
- Dickensheets D., and Steele A. (2002) Astrobiological instrumentation for Mars–the only
  way is down. *International Journal of Astrobiology*, 1: 365-380.
- Emerson D., Weiss J. V., and Megonigal J. P. (1999) Iron-oxidizing bacteria are associated with
- ferric hydroxide precipitates (Fe-plaque) on the roots of wetland plants. *Applied and Environmental Microbiology*, 65: 2758-2761.
- Fairén A. G., Parro V., Schulze-Makuch D., and Whyte L. (2017) Searching for life on Mars
  before it is too late. *Astrobiology*, 17: 962-970.
- Feng X., Matranga C., Vidic R., and Borguet E. (2004) A vibrational spectroscopic study of the
- fate of oxygen-containing functional groups and trapped CO2 in single-walled carbon
- nanotubes during thermal treatment. *The Journal of Physical Chemistry B*, 108: 19949-
- 770 19954.
- Frawley E. R., and Fang F. C. (2014) The ins and outs of bacterial iron metabolism. *Molecular microbiology*, 93: 609-616.
- Giovannetti R. (2012) The use of Spectrophotometry UV-Vis for the Study of Porphyrins. In:
   *Macro To Nano Spectroscopy*, InTech.
- 775 Glavin D. P., Freissinet C., Miller K. E., Eigenbrode J. L., Brunner A. E., Buch A., Sutter B.,
- Archer P. D., Atreya S. K., and Brinckerhoff W. B. (2013) Evidence for perchlorates and
- the origin of chlorinated hydrocarbons detected by SAM at the Rocknest aeolian deposit
- in Gale Crater. *Journal of Geophysical Research: Planets*, 118: 1955-1973.

779	Goordial J., Altshuler I., Hindson K., Chan-Yam K., Marcolefas E., and Whyte L. (2017) In situ
780	field sequencing and life detection in remote (79° 26' N) Canadian High Arctic
781	permafrost ice wedge microbial communities. Frontiers in microbiology, 8: 2594.
782	Goordial J., Davila A., Lacelle D., Pollard W., Marinova M. M., Greer C. W., DiRuggiero J.,
783	McKay C. P., and Whyte L. G. (2016) Nearing the cold-arid limits of microbial life in
784	permafrost of an upper dry valley, Antarctica. The ISME journal, 10: 1613.
785	Guinness E. A., Arvidson R. E., Clark I. H., and Shepard M. K. (1997) Optical scattering
786	properties of terrestrial varnished basalts compared with rocks and soils at the Viking
787	Lander sites. Journal of Geophysical Research: Planets, 102: 28687-28703.
788	Hays L. E., Graham H. V., Des Marais D. J., Hausrath E. M., Horgan B., McCollom T. M.,
789	Parenteau M. N., Potter-McIntyre S. L., Williams A. J., and Lynch K. L. (2017)
790	Biosignature preservation and detection in Mars analog environments. Astrobiology, 17:
791	363-400.
792	He Z., Gentry T. J., Schadt C. W., Wu L., Liebich J., Chong S. C., Huang Z., Wu W., Gu B., and
793	Jardine P. (2007) GeoChip: a comprehensive microarray for investigating
794	biogeochemical, ecological and environmental processes. The ISME journal, 1: 67.
795	Hecht M., Kounaves S., Quinn R., West S., Young S., Ming D., Catling D., Clark B., Boynton
796	W., and Hoffman J. (2009) Detection of perchlorate and the soluble chemistry of martian
797	soil at the Phoenix lander site. Science, 325: 64-67.
798	Jain M., Tyson J. R., Loose M., Ip C. L., Eccles D. A., O'Grady J., Malla S., Leggett R. M.,
799	Wallerman O., and Jansen H. J. (2017) MinION Analysis and Reference Consortium:
800	Phase 2 data release and analysis of R9. 0 chemistry. F1000Research, 6.

- 801 Jehlička J., Edwards H., and Oren A. (2013) Bacterioruberin and salinixanthin carotenoids of 802 extremely halophilic Archaea and Bacteria: a Raman spectroscopic study. Spectrochimica 803 Acta Part A: Molecular and Biomolecular Spectroscopy, 106: 99-103. 804 Jehlicka J., and Oren A. (2013) Raman spectroscopy in halophile research. Frontiers in 805 microbiology, 4: 380. 806 John K., Botkin D., Burton A., Castro-Wallace S., Chaput J., Dworkin J., Lehman N., Lupisella 807 M., Mason C., and Smith D. (2016) The Biomolecule Sequencer Project: Nanopore 808 sequencing as a dual-use tool for crew health and astrobiology investigations. 809 Johnson S. S., Zaikova E., Goerlitz D. S., Bai Y., and Tighe S. W. (2017) Real-Time DNA 810 Sequencing in the Antarctic Dry Valleys Using the Oxford Nanopore Sequencer. Journal 811 of Biomolecular Techniques: JBT, 28: 2. 812 Jones M. M., Jones E. A., Harmon D. F., and Semmes R. T. (1961) A search for perchlorate 813 complexes. Raman spectra of perchlorate solutions. Journal of the American Chemical 814 Society, 83: 2038-2042. 815 Kaczor A., and Baranska M. (2011) Structural changes of carotenoid astaxanthin in a single algal 816 cell monitored in situ by Raman spectroscopy. Analytical chemistry, 83: 7763-7770.
  - 817 Kim S. K., Kim M. S., and Suh S. W. (1987) Surface-enhanced Raman scattering (SERS) of

818 aromatic amino acids and their glycyl dipeptides in silver sol. Journal of Raman

- 819 *spectroscopy*, 18: 171-175.
- Koren S., Walenz B. P., Berlin K., Miller J. R., Bergman N. H., and Phillippy A. M. (2017)
- 821 Canu: scalable and accurate long-read assembly via adaptive k-mer weighting and repeat
  822 separation. *Genome research*, 27: 722-736.

- 823 Kozich J. J., Westcott S. L., Baxter N. T., Highlander S. K., and Schloss P. D. (2013)
- 824 Development of a dual-index sequencing strategy and curation pipeline for analyzing
- 825 amplicon sequence data on the MiSeq Illumina sequencing platform. *Applied and*
- 826 *environmental microbiology*: AEM. 01043-13.
- 827 Lacap D. C., Warren-Rhodes K. A., McKay C. P., and Pointing S. B. (2011) Cyanobacteria and
- 828 chloroflexi-dominated hypolithic colonization of quartz at the hyper-arid core of the
  829 Atacama Desert, Chile. *Extremophiles*, 15: 31-38.
- 830 Lee K. C.-Y., Herbold C., Dunfield P. F., Morgan X. C., McDonald I. R., and Stott M. B. (2013)
- 831 Phylogenetic delineation of the novel phylum Armatimonadetes (former candidate
- division OP10) and definition of two novel candidate divisions. *Appl. Environ*.
- 833 *Microbiol.*, 79: 2484-2487.
- Leggett R. M., and Clark M. D. (2017) A world of opportunities with nanopore sequencing. *Journal of Experimental Botany*.
- Loman N. J., and Watson M. (2015) Successful test launch for nanopore sequencing. *Nature methods*, 12: 303-304.
- 838 Luef B., Fakra S. C., Csencsits R., Wrighton K. C., Williams K. H., Wilkins M. J., Downing K.
- H., Long P. E., Comolli L. R., and Banfield J. F. (2013) Iron-reducing bacteria
- 840 accumulate ferric oxyhydroxide nanoparticle aggregates that may support planktonic
- growth. *The ISME journal*, 7: 338.
- 842 Maiti N. C., Apetri M. M., Zagorski M. G., Carey P. R., and Anderson V. E. (2004) Raman
- 843 spectroscopic characterization of secondary structure in natively unfolded proteins: α-
- synuclein. *Journal of the American Chemical Society*, 126: 2399-2408.

- Manley L. J., Ma D., and Levine S. S. (2016) Monitoring error rates in Illumina sequencing. *Journal of biomolecular techniques: JBT*, 27: 125.
- 847 Markowitz V. M., Chen I.-M. A., Palaniappan K., Chu K., Szeto E., Grechkin Y., Ratner A.,
- 848Jacob B., Huang J., and Williams P. (2011) IMG: the integrated microbial genomes
- database and comparative analysis system. *Nucleic acids research*, 40: D115-D122.
- 850 Marshall C. P., Carter E. A., Leuko S., and Javaux E. J. (2006) Vibrational spectroscopy of
- extant and fossil microbes: relevance for the astrobiological exploration of Mars.
- 852 *Vibrational Spectroscopy*, 41: 182-189.
- 853 Martins Z., Sephton M., Foing B., and Ehrenfreund P. (2011) Extraction of amino acids from
- soils close to the Mars Desert Research Station (MDRS), Utah. *International Journal of Astrobiology*, 10: 231-238.
- 856 McCauley J., Breed C., El-Baz F., Whitney M., Grolier M., and Ward A. (1979) Pitted and fluted
- 857 rocks in the Western Desert of Egypt: Viking comparisons. *Journal of Geophysical*858 *Research: Solid Earth*, 84: 8222-8232.
- McLennan S., Sephton M., Beaty D., Hecht M., Pepin B., Leya I., Jones J., Weiss B., Race M.,
- and Rummel J. (2012) Planning for Mars returned sample science: final report of the
- MSR End-to-End International Science Analysis Group (E2E-iSAG). *Astrobiology*, 12:
  175-230.
- 863 Meinert C., Myrgorodska I., De Marcellus P., Buhse T., Nahon L., Hoffmann S. V.,
- d'Hendecourt L. L. S., and Meierhenrich U. J. (2016) Ribose and related sugars from
- ultraviolet irradiation of interstellar ice analogs. *Science*, 352: 208-212.

- 866 Menegon M., Cantaloni C., Rodriguez-Prieto A., Centomo C., Abdelfattah A., Rossato M.,
- Bernardi M., Xumerle L., Loader S., and Delledonne M. (2017) On site DNA barcoding
  by nanopore sequencing. *PloS one*, 12: e0184741.
- 869 Mileikowsky C., Cucinotta F. A., Wilson J. W., Gladman B., Horneck G., Lindegren L., Melosh
- 870 J., Rickman H., Valtonen M., and Zheng J. (2000) Natural transfer of viable microbes in
- space: 1. From Mars to Earth and Earth to Mars. *Icarus*, 145: 391-427.
- 872 Musk E. (2017) Making humans a multi-planetary species. *New Space*, 5: 46-61.
- 873 Mustard J., Adler M., Allwood A., Bass D., Beaty D., and Bell J. (2013a) Appendices to the
- 874 Report of the Mars 2020 Science Definition Team. *Mars Exploration Program Analysis*875 *Group*: 154.
- 876 Mustard J., Adler M., Allwood A., Bass D., Beaty D., Bell J., Brinckerhoff W., Carr M., Des

877 Marais D., and Brake B. (2013b) Report of the mars 2020 science definition team. *Mars*878 *Explor. Progr. Anal. Gr*: 155-205.

- Neveu M., Hays L. E., Voytek M. A., New M. H., and Schulte M. D. (2018) The Ladder of Life
  Detection. *Astrobiology*.
- 881 Niederberger T. D., Perreault N. N., Tille S., Lollar B. S., Lacrampe-Couloume G., Andersen D.,

882 Greer C. W., Pollard W., and Whyte L. G. (2010) Microbial characterization of a

- subzero, hypersaline methane seep in the Canadian High Arctic. *The ISME journal*, 4:
- 884 1326.
- NRC. (2012) Vision and voyages for planetary science in the decade 2013-2022. National
  Academies Press.

- 887 Oren A., Sørensen K. B., Canfield D. E., Teske A. P., Ionescu D., Lipski A., and Altendorf K.
- 888 (2009) Microbial communities and processes within a hypersaline gypsum crust in a
- saltern evaporation pond (Eilat, Israel). *Hydrobiologia*, 626: 15-26.
- 890 Osinski G. R., Battler M., Caudill C. M., Francis R., Haltigin T., Hipkin V. J., Kerrigan M.,
- 891 Pilles E., Pontefract A., and Tornabene L. L. (2018) The CanMars Mars Sample Return
  892 analogue mission. *Planetary and Space Science*, (in press).
- Pace N. R. (2001) The universal nature of biochemistry. *Proceedings of the National Academy of Sciences*, 98: 805-808.
- 895 Parro V., de Diego-Castilla G., Rodríguez-Manfredi J. A., Rivas L. A., Blanco-López Y.,
- 896 Sebastián E., Romeral J., Compostizo C., Herrero P. L., and García-Marín A. (2011)
- 897 SOLID3: a multiplex antibody microarray-based optical sensor instrument for in situ life
  898 detection in planetary exploration. *Astrobiology*, 11: 15-28.
- 899 Parro V., Rodríguez-Manfredi J., Briones C., Compostizo C., Herrero P., Vez E., Sebastián E.,
- 900 Moreno-Paz M., García-Villadangos M., and Fernández-Calvo P. (2005) Instrument
- 901 development to search for biomarkers on Mars: terrestrial acidophile, iron-powered
- 902 chemolithoautotrophic communities as model systems. *Planetary and Space Science*, 53:
  903 729-737.
- 904 Pontefract A., Hachey J., Zuber M. T., Ruvkun G., and Carr C. E. (2018) Sequencing Nothing:
- 905 Exploring Failure Modes of Nanopore Sensing and Implications for Life Detection. *Life*906 Sciences in Space Research.
- 907 Poretsky R., Rodriguez-R L. M., Luo C., Tsementzi D., and Konstantinidis K. T. (2014)
- 908 Strengths and limitations of 16S rRNA gene amplicon sequencing in revealing temporal
- 909 microbial community dynamics. *PloS one*, 9: e93827.

- Raes J., Korbel J. O., Lercher M. J., Von Mering C., and Bork P. (2007) Prediction of effective
  genome size in metagenomic samples. *Genome biology*, 8: R10.
- 912 Ranjan R., Rani A., Metwally A., McGee H. S., and Perkins D. L. (2016) Analysis of the
- 913 microbiome: Advantages of whole genome shotgun versus 16S amplicon sequencing.
- 914 Biochemical and biophysical research communications, 469: 967-977.
- 915 Rhind T., Ronholm J., Berg B., Mann P., Applin D., Stromberg J., Sharma R., Whyte L., and
- 916 Cloutis E. (2014) Gypsum-hosted endolithic communities of the Lake St. Martin impact
- 917 structure, Manitoba, Canada: spectroscopic detectability and implications for Mars.
- 918 International Journal of Astrobiology, 13: 366-377.
- 919 Rivas L. A., Aguirre J., Blanco Y., González-Toril E., and Parro V. (2011) Graph-based
- deconvolution analysis of multiplex sandwich microarray immunoassays: applications for
  environmental monitoring. *Environmental microbiology*, 13: 1421-1432.
- 922 Rivas L. A., García-Villadangos M., Moreno-Paz M., Cruz-Gil P., Gómez-Elvira J., and Parro V.
- 923 (2008) A 200-antibody microarray biochip for environmental monitoring: searching for
- 924 universal microbial biomarkers through immunoprofiling. *Analytical chemistry*, 80:
- 925 7970-7979.
- 926 Ruan C., Wang W., and Gu B. (2006) Surface-enhanced Raman scattering for perchlorate

927 detection using cystamine-modified gold nanoparticles. *Analytica chimica acta*, 567:

- 928 114-120.
- 929 Rull F., Maurice S., Hutchinson I., Moral A., Perez C., Diaz C., Colombo M., Belenguer T.,
- 930 Lopez-Reyes G., and Sansano A. (2017) The Raman Laser Spectrometer for the ExoMars
  931 Rover Mission to Mars. *Astrobiology*, 17: 627-654.

932	Sánchez-García L.,	Fernández M.,	García-Villadang	os M.,	Blanco Y.	, Cady S.	, Hinman N.,
	,		· · · · · · · · · · · · · · · · · · ·	, ,	,	, J	/ /

- Bowden M., Pointing S., Lee K., Warren-Rhodes K. and others. (2018) Microbial
- biomarker transition in high altitude sinter mounds from El Tatio (Chile) through
- 935 different stages of hydrothermal activity. *Frontiers in microbiology*, (under review).
- Sandford S., and Allamandola L. (1990) The physical and infrared spectral properties of CO2 in
  astrophysical ice analogs. *The Astrophysical Journal*, 355: 357-372.
- 938 Schlesinger W. H., Pippen J. S., Wallenstein M. D., Hofmockel K. S., Klepeis D. M., and Mahall
- B. E. (2003) Community composition and photosynthesis by photoautotrophs under

940 quartz pebbles, southern Mojave Desert. *Ecology*, 84: 3222-3231.

- 941 Sephton M. A. (2018) Selecting Mars samples to return to Earth. *Astronomy & Geophysics*, 59:
  942 1.36-1.38.
- 943 Spry J. A., Race M., Kminek G., Siegel B., and Conley C. (2018) Planetary Protection
- 944 Knowledge Gaps for Future Mars Human Missions: Stepwise Progress in Identifying and
- 945 Integrating Science and Technology Needs. 48th International Conference on
- 946 Environmental Systems.
- 947 Stoker C. R., Clarke J., Direito S. O., Blake D., Martin K. R., Zavaleta J., and Foing B. (2011)

948 Mineralogical, chemical, organic and microbial properties of subsurface soil cores from

- 949 Mars Desert Research Station (Utah, USA): Phyllosilicate and sulfate analogues to Mars
- 950 mission landing sites. *International Journal of Astrobiology*, 10: 269-289.

951 Stromberg J., Applin D., Cloutis E., Rice M., Berard G., and Mann P. (2014) The persistence of

- a chlorophyll spectral biosignature from Martian evaporite and spring analogues under
- 953 Mars-like conditions. *International Journal of Astrobiology*, 13: 203-223.

- 954 Summons R., Sessions A., Allwood A., Barton H., Beaty D., Blakkolb B., Canham J., Clark B.,
- 955 Dworkin J., and Lin Y. (2014) Planning considerations related to the organic
- 956 contamination of Martian samples and implications for the Mars 2020 rover.
- 957 *Astrobiology*, 14: 969-1027.
- Tahon G., Tytgat B., Lebbe L., Carlier A., and Willems A. (2018) Abditibacterium utsteinense
- 959 sp. nov., the first cultivated member of candidate phylum FBP, isolated from ice-free
  960 Antarctic soil samples. *Systematic and applied microbiology*, 41: 279-290.
- 961 Tessler M., Neumann J. S., Afshinnekoo E., Pineda M., Hersch R., Velho L. F. M., Segovia B.
- 962 T., Lansac-Toha F. A., Lemke M., and DeSalle R. (2017) Large-scale differences in
- 963 microbial biodiversity discovery between 16S amplicon and shotgun sequencing.
- 964 *Scientific reports*, 7: 6589.
- Thiel C. S., Ehrenfreund P., Foing B., Pletser V., and Ullrich O. (2011) PCR-based analysis of
   microbial communities during the EuroGeoMars campaign at Mars Desert Research

967 Station, Utah. *International Journal of Astrobiology*, 10: 177-190.

- 968 Thomas G. J., Prescott B., and Urry D. W. (1987) Raman amide bands of type-II β-turns in
- 969 cyclo-(VPGVG) 3 and poly-(VPGVG), and implications for protein secondary-structure
  970 analysis. *Biopolymers*, 26: 921-934.
- 971 Van Lith Y., Warthmann R., Vasconcelos C., and Mckenzie J. A. (2003) Sulphate-reducing
- 972 bacteria induce low-temperature Ca-dolomite and high Mg-calcite formation.
- 973 *Geobiology*, 1: 71-79.
- Warren-Rhodes K. A., Rhodes K. L., Pointing S. B., Ewing S. A., Lacap D. C., Gomez-Silva B.,
  Amundson R., Friedmann E. I., and McKay C. P. (2006) Hypolithic cyanobacteria, dry

- 976 limit of photosynthesis, and microbial ecology in the hyperarid Atacama Desert.
- 977 *Microbial ecology*, 52: 389-398.
- 978 Williams R. M., Chidsey Jr T. C., and Eby D. E. (2007) Exhumed paleochannels in central
- 979 Utah—Analogs for raised curvilinear features on Mars.
- 980 Williams R. M., Irwin III R. P., Burr D. M., Harrison T., and McClelland P. (2013) Variability in
- 981 Martian sinuous ridge form: Case study of Aeolis Serpens in the Aeolis Dorsa, Mars, and 982 insight from the Mirackina paleoriver, South Australia. *Icarus*, 225: 308-324.
- 983 Williams R. M., Irwin R. P., and Zimbelman J. R. (2009) Evaluation of paleohydrologic models
- 984 for terrestrial inverted channels: Implications for application to martian sinuous ridges.
- 985 *Geomorphology*, 107: 300-315.
- 286 Ziolkowski L., Mykytczuk N., Omelon C., Johnson H., Whyte L., and Slater G. (2013) Arctic
- 987 gypsum endoliths: a biogeochemical characterization of a viable and active microbial
- 988 community.
- 989
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999 Tables

- 1001 Table 1. Input DNA and MinION sequence details from the Hanksville paleochannel samples.
- 1002 Also included are the number of correct reads by Canu, the percent of reads classified by JGI
- 1003 IMG/M ER, and the JGI IMG/M ER submission ID.

Sample type	Sample name	Input DNA (ng)	Total reads	Avg. sequence length (bp)	Corrected reads	% of hits classified	JGI Submission ID
Endolith	E1	5.07	1971	2939	589	2.54	184799
Endolith	E2	4.5	4026	3056	1683	2.61	184800
Endolith	E3	4.71	11 943	3372	8261	7.96	185540
Endolith	E4	4.04	9885	3571	6643	9.15	185543
Endolith	E5	4.475	9189	3513	6317	8.43	185546
Regolith	<b>S</b> 1	0.756	2446	2598	868	2.64	185544
Regolith	S2	0.667	3580	3293	1519	0.96	185545
Regolith	<b>S</b> 3	9	3352	3021	1434	1.22	185548
Rock	R1	3.03	4702	2884	2654	6.77	185547
Rock	R2	19.5	1594	2920	369	0.72	185559
Rock	R3	3.975	8461	3310	4560	4.46	185550

- 1014 Table 2. DNA detection limit trial details with endolith sample E4. Also included are the number
- 1015 of correct reads by Canu, the percent of reads classified by JGI IMG/M ER, and the JGI IMG/M
- 1016 ER submission ID.

	Sample	Input DNA (ng)	Total reads	Corrected reads	% of hits classified	JGI Submission ID
		0.110	253 188	109 064	31.24	190088
	Endolith E4	0.0408	192 959	65 445	5.89	195531
	Elluollul E4	0.0106	64 300	21 582	1.56	195532
		0.00106	1448	513	2.14	195533
	No DNA (3					
	µl nuclease-	0	121	55	0	190089
	free water)					
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Table 3. Relationship between MinION sequences, XRD data, reflectance spectra, Raman spectra, and LDChip immunoassays. A feature is presented in the left column, and its presence, potential presence as microbial taxa, or resulting trait in the corresponding techniques is noted.

Facture	Technique						
reature	XRD	Reflectance	Raman	MinION	LDChip		
Quartz	Major	N/A	470 cm <sup>-1</sup> peak	Cyanobacteria	Cyanobacteria		
Calcite	Minor	N/A	N/A	Cyanobacteria, SRBs, <i>Bacillus</i> , sulfate system/sulfate-sulfur system transport genes, sulfur metabolism	Cyanobacteria, Bacillus, SRBs		
Gypsum	Minor	N/A	1006 cm <sup>-1</sup> peak	Cyanobacteria, Halobacteria, SRBs, sulfate system/sulfate- sulfur system transport genes, sulfur metabolism	Cyanobacteria, Haloferax, SRBs		
Iron	N/A	Fe-bearing oxyhydroxides (~ 500 nm, ~ 900 nm, ~ 1150 nm), Fe-bearing mineral charge transfers (> ~ 1000 nm)	Background Fe fluorescence	Heme oxygenase	Iron reductase, iron/sulfur environmental samples from Río Tinto		
Perchlorate	N/A	N/A	~ 930 cm <sup>-1</sup>	N/A	Perchlorate reductase		

Chlorophyll	N/A	670 nm, chlorophyll bump, red edge, porphyrins at ~ 400 nm	~ 3530 cm <sup>-1</sup>	Cyanobacteria, chlorophyll photosystem I and II genes, protochlorophyllide reductase, protoporphyrin IX cyclase	Cyanobacteria
Carotenoids	N/A	N/A	N/A	Cyanobacteria, β-carotene ketolase and hydroxylase	Cyanobacteria
Nitrogen metabolism	N/A	N/A	N/A	NifU, NirS	NifD1, nitrate reductase
Methane metabolism	N/A	N/A	N/A	Methanobacteria, Methanomicrobia	Methanobacterium

# **Figure Legends**

**FIG. 1.** Hanksville paleochannel site and samples. (**A**) Side view of Jotunheim (38.41709915, -110.78539335), an inverted river channel. (**B**) Aerial view of the study site. Two main inverted channels are circled; Jotunheim is circled in the black, solid line. (**C**) Regolith sample S2 *in situ* during collection. Surface regolith is dark red and visibly desiccated, and covers a layer of light green sand.

**FIG. 2.** Metagenomes of the Hanksville paleochannel samples generated by MinION sequencing. All three domains of life are detected, as well as viruses. Samples E1, E2, E3, E4, and E5 are endoliths. Samples S1, S2, and S3 are regolith. Samples R1, R2, and R3 are non-endolith rocks.

FIG. 3. (A) Bacterial 16S rRNA community profile of the Hanksville paleochannels generated

by Illumina MiSeq sequencing. (B) Bacterial content of the whole shotgun metagenome from

MinION sequencing of the Hanksville paleochannels.

**FIG. 4.** Weighted UniFrac principal coordinates analysis (PCoA) of the 16S rRNA sequencing data of the Hanksville paleochannels. Samples are coloured by type: endoliths in green (E1, E2, E3, E4, and E5), sediment in red (S1, S2, and S3), and non-endolith rocks in blue (R1, R2, and

R3). The variation between samples is expressed as percentages along each axis.

FIG. 5. Detection limit trials with endolith sample E4 at the phyla level. The first column "E4" is

the same data as presented in Figure 2. Taxonomy and abundance change with input DNA, and

the final 0.001 ng trial bears little resemblance to E4's original community composition.

**FIG. 6.** Reflectance spectra of representative samples from the Hanksville paleochannels with major absorption features labelled. (**A**) Endolith-bearing samples have a prominent feature at 670 nm indicative of chlorophyll. The dark surface of the sample also shows a minor feature at 670 nm which may indicate desiccated endolithic communities. (**B**) Spectra of non-endolith bearing samples varies as a reflection of mineralogy. The red color of regolith sample S1 is reflected in the strong Fe absorptions in the VNIR region of the spectrum, and also shows evidence for carbonate bearing phases.

**FIG. 7.** Raman spectra of representative samples from the Hanksville palechannels with major features labelled. (**A**) Endolith-bearing samples have spectral features consistent with the presence of chlorophyll and other organic compounds. (**B**) The Raman spectra of non-endolith bearing samples are dominated by a broad Fe flourescence feature.

**FIG. 8.** Heat map comparing positive LDChip and MinION results at the phylum level. Relative abundance, which represents the value of signal intensity corresponding to each positive antibody within the category with regards to the other categories and to the total signal intensity in the sample, was plotted in a scale from white (negative results) to red (maximum). The maximum value was 0.53 for LDChip results. Total number of MinION sequences corresponding to the listed antibodies were counted and are presented on a scale from 0 - 1.

**FIG. 9.** Heat map comparing positive LDChip and MinION results for specific protein detection. Relative abundance, which represents the value of signal intensity corresponding to each positive antibody within the category with regards to the other categories and to the total signal intensity in the sample, was plotted in a scale from white (negative results) to red (maximum). The maximum value was 0.53 for LDChip results. Total number of MinION sequences corresponding to the listed antibodies were counted and are presented on a scale from 0 - 1.