GEOCHEMISTRY

Minerals with metal-organic framework structures

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Metal-organic frameworks (MOFs) are an increasingly important family of advanced materials based on open, nanometer-scale metal-organic architectures, whose design and synthesis are based on the directed assembly of carefully designed subunits. We now demonstrate an unexpected link between mineralogy and MOF chemistry by discovering that the rare organic minerals stepanovite and zhemchuzhnikovite exhibit structures found in well-established magnetic and proton-conducting metal oxalate MOFs. Structures of stepanovite and zhemchuzhnikovite, exhibiting almost nanometer-wide and guest-filled apertures and channels, respectively, change the perspective of MOFs as exclusively artificial materials and represent, so far, unique examples of open framework architectures in organic minerals.

INTRODUCTION

Metal-organic frameworks (MOFs) are a rapidly developing family of advanced materials, with applications in hydrogen storage, carbon sequestration, catalysis, gas separation, photovoltaics, and more (1-4), synthesized from inorganic and organic components judiciously selected to form two- or three-dimensional architectures with open or potential voids (5-7). Here, we demonstrate the appearance of MOF structures in nature as rare oxalate minerals stepanovite and zhemchuzhnikovite (8, 9) that are based on open anionic frameworks analogous to those found in magnetic (10) or proton-conductive (11) metal oxalate MOFs. Metal oxalates are the largest family of organic minerals (12), mainly generated by biogeochemical processes such as mineral weathering and in vivo formation by living organisms (13, 14). Their structures are dominated by mononuclear complexes and one-dimensional coordination polymers of hydrated metal ions [Mg²⁺ in glushinskite (15), Cu²⁺ in moolooite (16), Mn²⁺ in lindbergite (17), and Fe^{2+} in humboldtine (18)] bridged by oxalate anions (\mathbf{ox}^{2-}) . However, there are several oxalate minerals of more complex compositions that have not yet been structurally characterized, notably stepanovite (Strunz class 10.AB.20, reported in 1942) and zhemchuzhnikovite (Strunz class 10.AB.35, reported in 1963), first discovered near the estuary of Lena River (Sakha-Yakutia, Siberia, Russia) and accepted by the International Mineralogical Association in 1964 (8, 9, 12). Both minerals appear as thin veinlets, which consist of green transparent grained or fibrous aggregates with, very rarely, isometric crystals ~0.05 mm across (Fig. 1A and fig. S1), in lignite. They are associated with other natural salts of organic acids, for example, oxalates (whewellite, weddellite, and glushinskite) and unspecified acetates, and with calcite and dolomite. The rarity of stepanovite and zhemchuzhnikovite in nature is definitely caused by very uncommon geochemical environment and physical conditions of their formation: at both coal deposits Tyllakh and Chai-Tumus, these water-soluble minerals were found in drill cores at depths up to 230 m below the day surface, in the permafrost zone, within lignite saturated by natural acetic acid (8). The formulae of stepanovite $[NaMgFe(\mathbf{ox})_3 \cdot 8-9H_2O]$

and zhemchuzhnikovite $[NaMg(Fe_{0.4}Al_{0.6})(\mathbf{ox})_3\cdot 8-9H_2O]$ were reported by Knipovich *et al.* (8), along with crystallographic parameters (Table 1) that were determined from powder x-ray diffraction (PXRD) data. We noted that the mineral compositions of stepanovite and zhemchuzhnikovite closely resemble those of anionic oxalate MOFs first developed in the 1990s as magnetic materials (10). These MOFs consist of open $[M^IM^{III}(\mathbf{ox})_3]^{2-}$ frameworks of three-dimensional (3D) (10,3) net or two-dimensional (2D) (6,3) net topology [honeycomb (**hcb**) topology; Fig. 1B], with monovalent (M^I ; for example, Li^+ or Na^+) and trivalent cations (M^{III} ; for example, Cr^{3+} or Fe^{3+}) as nodes, and are templated by divalent cations retained in framework cavities. More recently, analogous metal-oxalate structures based on zinc (19, 20) or other transition metals (21) have garnered additional interest as ferromagnetic and/or proton-conducting materials (Fig. 1C) (11, 21).

The unexpected chemical similarity to known MOFs led us to speculate whether stepanovite and zhemchuzhnikovite might be based on similar open frameworks, prompting structural characterization. Searching the Cambridge Structural Database (CSD) for structures with crystallographic parameters (Table 1) resembling those for stepanovite gave no meaningful results. However, a search for zhemchuzhnikovite parameters (Table 1) revealed isostructurality to NaMgAl(**ox**)₃·9H₂O [Cambridge Crystallographic Data Centre (CCDC) code YODWUK] (*22*), a 2D MOF consisting of **hcb** [NaAl(**ox**)₃^{2–} open framework layers, with Mg(H₂O)₆²⁺ guests in layer cavities. Isostructurality and chemical similarity to NaMgAl(**ox**)₃·9H₂O reinforced our view that zhemchuzhnikovite exhibits an MOF structure.

RESULTS

For this study, we used samples of minerals collected by the Russian mineralogist P. I. Glushinsky in 1942, later studied by Knipovich *et al.* (8). Because natural samples appear together with other minerals, we also pursued the preparation of synthetic stepanovite and zhemchuzhnikovite to obtain phase-pure materials and single crystals of improved quality. Synthetic stepanovite was obtained by reacting Fe_2O_3 and MgO with aqueous NaOH and oxalic acid (23). After 2 days, the solution yielded green crystals of trigonal habit, consistent with that of stepanovite (Fig. 1D and fig. S2) (8, 9, 12). Synthetic zhemchuzhnikovite was made

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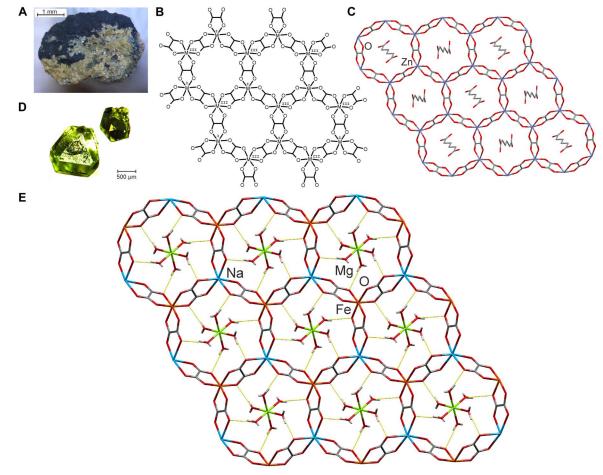


Fig. 1. Stepanovite is a mineral with an MOF structure. (**A**) Stepanovite sample, Chai-Tumus coal deposit (Sakha-Yakutia, Siberia, Russia; sample from E. I. Nefedov's collection). (**B**) schematic of an open anionic **hcb** framework composed of M^{I} and M^{III} nodes bridged by oxalates. (**C**) A single layer of an analogous zinc-based proton-conducting MOF material, including guests (*11*). (**D**) Crystals of synthetic stepanovite. (**E**) A single metal-organic layer in stepanovite, viewed along the crystallographic *c* axis, displaying the anionic **hcb** $[NaFe(ox)_3]^{2-}$ framework, with apertures occupied by Mg(H₂O)₆²⁺. Hydrogen bonds between Mg(H₂O)₆²⁺ guests and $[NaFe(ox)_3]^{2-}$ framework are highlighted as yellow dotted lines.

by reacting premade $Na_3Fe(\mathbf{ox})_3$ ·3H₂O, aluminum sulfate, and MgO with oxalic acid in water, using a 6:4 stoichiometric ratio of Al and Fe, respectively. The reaction yielded pale yellow-green crystals, with morphology consistent with that of zhemchuzhnikovite (fig. S3).

Single-crystal X-ray diffraction (XRD) structures of synthetic and natural stepanovite and zhemchuzhnikovite confirmed the unit-cell parameters reported by Knipovich *et al.* (8) (Table 1) and gave unambiguous evidence that these minerals exhibit structures previously found in 2D oxalate MOFs. The stepanovite structure consists of anionic **hcb** $[NaFe(\mathbf{ox})_3]^{2-}$ layers, each layer exhibiting apertures of ~0.9 nm in diameter, occupied with Mg(H₂O)₆²⁺ guests hydrogen-bonded to oxalate ions of the pore wall (O···O distances, 2.82 to 2.96 Å; O–H···O angles, 148° to 167°) (Fig. 1E). In each layer, Fe³⁺ and Na⁺ adopt octahedral environments of opposite chirality (Λ or Δ).

Layers are separated by water molecules, each forming four hydrogen bonds: two as donors toward oxalates of neighboring layers (O···O distances, 2.74 to 2.77 Å; O–H···O angles, 168° to 175°) and two as hydrogen bond acceptors from $Mg(H_2O)_6^{2+}$ sitting in neighboring layers (O···O distances, 2.76 to 2.79 Å; O–H···O angles, 165° to 176°). Hydrogen bonds around each water molecule form a tetrahedron elongated in crystallographic c direction (Fig. 2A).

Crystal structure determination and refinement of natural and synthetic zhemchuzhnikovite revealed open hcb topology layers with Al^{3+} and Fe^{3+} disordered over the same site (Fig. 3A). Depending on crystal choice, occupational parameters of Fe³⁺ and Al³⁺ in synthetic crystals varied between 0.41:0.59 and 0.76:0.24, whereas a single crystal from the natural sample gave 0.69:0.31. For each investigated crystal, the Fe³⁺/Al³⁺ disorder appeared statistical, without no indication of ordering into a supercell. We conclude that the structure of zhemchuzhnikovite, NaMg(Fe_{1-x}Al_x)(\mathbf{ox})₃·9H₂O, can accommodate diverse Al/Fe compositions, also evidenced by atomic emission analysis of single crystals from different batches of synthetic zhemchuzhnikovite, with x in the range of 0.21 to 0.86. The structure of NaMg(Fe_{0.41}Al_{0.59}) $(\mathbf{ox})_3$ ·9H₂O was anisotropically refined to $R_1 < 0.04$ in space group P3c1, identical to NaMgAl(ox)3.9H2O (22). Attempted refinement in centrosymmetric group $P\overline{3}c1$ gave a poorer result, that is, $R_1 > 0.08$. As in stepanovite, cavities of each layer are occupied by Mg(H₂O)₆²⁺ hydrogenbonded to oxalates lining the pore [O…O distances, 2.83 to 2.99 Å;

	:	Stepanovite	
	Reported by Knipovich et al. (8)	Natural	Synthetic
Crystal system	Trigonal	Trigonal	Trigonal
Formula	NaMgFe(C ₂ O ₄) ₃ ·8-9H ₂ O	NaMgFe(C ₂ O ₄) ₃ ·9H ₂ O	NaMgFe(C ₂ O ₄) ₃ ·9H ₂ O
Space group	Not reported	R3c	R3c
a (Å)	9.78 [*]	9.8367(13)*	9.887(13)*
c (Å)	36.67	36.902(5)	37.03(5)
a/c ratio	1:3.73–1:3.76*	1:3.75*	1:3.75*
V (Å ³)	3070	3092.2(7)	3135(9)
Ζ	6	6	6
Color	Green	Greenish yellow	Green
Density (g cm ⁻³)	1.69	1.71 (calculated)	1.68 (calculated)
	Zhei	nchuzhnikovite	•
	Reported by Knipovich et al. (8)	Natural	Synthetic
Crystal system	Trigonal	Trigonal	Trigonal
Formula	NaMg(Fe _{0.4} Al _{0.6})(C ₂ O ₄) ₃ ·8-9H ₂ O	NaMg(Fe _{0.31} Al _{0.69})(C ₂ O ₄) ₃ ·9H ₂ O	NaMg(Fe _{1-x} Al _x)(C ₂ O ₄) ₃ ·9H ₂ O [†]
Space group	Not reported	P3c1 [‡]	P3c1 [‡]
a (Å)	16.67	16.809(7)	16.919(2) [§]
c (Å)	12.51	12.658(6)	12.561(2) [§]
a/c ratio	1:0.75–1:0.739	1:0.753	1:0.742 [§]
V (Å ³)	3001	3097(2)	3113.8(9) [§]
Ζ	6	6	6
Color	Green	Greenish yellow	Yellow-green
Density	1.62–1.66	1.64 (calculated)	1.63 (calculated)

Table 1. Crystallographic data for stepanovite and zhemchuzhnikovite. Comparison of crystallographic and general parameters reported by Knipovich *et al.* (8) to the investigated natural and synthetic samples in this study.

O–H…O angles, 161° to 177° for NaMg(Fe $_{0.41}$ Al $_{0.59}$)(**ox**) $_3$ ·9H₂O], and the layers are interconnected by hydrogen bonding to interstitial water.

In each **hcb** layer of zhemchuzhnikovite, all Fe³⁺ ions adopt the same chirality, opposite to that of Na⁺ in the same layer. However, unlike stepanovite, layers in zhemchuzhnikovite form ABABAB stacks with all **hcb** apertures aligned to form channels of ~0.9-nm diameter. The difference between stepanovite and zhemchuzhnikovite structures is evident in experimental and simulated PXRD patterns of zhemchuzhnikovite, which are noticeably different from those of stepanovite (Fig. 2C). The PXRD patterns of natural and synthetic zhemchuzhnikovite showed excellent fit to that which was calculated on the basis of the herein determined structure (table S2). The hydrogen bonding arrangement around interstitial water molecules in zhemchuzhnikovite resembles stepanovite: Each water molecule is a twofold donor toward oxalate ions in neighboring layers (O-O distances, 2.72 to 2.80 Å; O-H--O angles, 150° to 176°) and a twofold acceptor toward guest Mg(H₂O)₆²⁺ in neighboring layers [O--O distances, 2.74 to 2.75 Å;

O–H…O angles, 158° to 175° for NaMg(Fe_{0.41}Al_{0.59})(\mathbf{ox})₃·9H₂O]. Composition of synthetic zhemchuzhnikovite was confirmed by thermogravimetric analysis (TGA) (fig. S5).

The structural similarity of stepanovite and zhemchuzhnikovite to the proton-conducting oxalate MOFs is marked (11, 21). Proton conductivity in these MOFs results largely from a Grotthuss-type proton-hopping mechanism, enabled by a 2D network of hydrogen bonds involving water molecules in the interlayer space and protic species located either between the MOF layers or lodged in the pores (11, 24). On the basis of these considerations, stepanovite exhibits potential for proton conduction, because there is an uninterrupted 2D net of short hydrogen bonds in the interlayer space, involving interstitial water molecules and Mg(H₂O)₆²⁺ cations. Hydrogen bonding in both mineral structures extends beyond interlayer space, in the crystallographic *z* direction, giving a 3D network for stepanovite (Fig. 4A) and one-dimensional columns extending through MOF channels (Fig. 4B) for zhemchuzhnikovite.

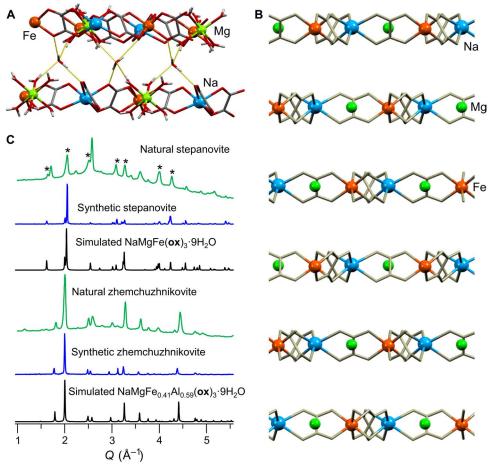


Fig. 2. Structure of stepanovite and PXRD patterns of MOF minerals. (**A**) Hydrogen bonding environment of water guests between **hcb** layers of stepanovite. (**B**) Stepanovite structure viewed parallel to crystallographic *a* axis (water molecules were omitted for clarity), with offset ABCABC arrangement evident from stacking of Na⁺, Mg²⁺, and Fe³⁺ ions (blue, green, and orange, respectively) in neighboring layers. (**C**) Comparison of PXRD patterns (top to bottom): natural stepanovite, synthetic stepanovite, simulated NaMgFe(**ox**)₃·9H₂O, natural zhemchuzhnikovite, synthetic zhemchuzhnikovite, and simulated NaMgFe_{0.41}Al_{0.59}(**ox**)₃·9H₂O. The PXRD pattern of natural stepanovite is affected by the presence of a number of other minerals, of which the most abundant one is glushinskite, as well as amorphous organic material. For clarity, principal reflections of stepanovite are designated with "*." List of indexed x-ray reflections for natural stepanovite and zhemchuzhnikovite is given in tables S1 and S2.

Thermal stability of zhemchuzhnikovite and stepanovite is consistent with the properties of proton-conducting oxalate MOFs reported by Sadakiyo et al. (11), with mild heating (90°C) leading to partial dehydration, which is reversed upon exposure to saturated water vapor at room temperature. TGA after thermal dehydration reveals that stepanovite loses 3 equiv of water upon heating, consistent with removal of water between the hcb layers. For zhemchuzhnikovite, heating leads to the loss of crystallinity and removal of 6 equiv of water, indicating the loss of at least 50% of coordinated water guests in MOF channels. Upon exposure to saturated water vapor at room temperature, both stepanovite and zhemchuzhnikovite regain the original structure and water content (figs. S6 to S10). The removal and absorption of water can be repeated at least three times (Fig. 4C and fig. S6) for both minerals, each time regenerating the initial MOF. The fact that stepanovite loses only interstitial water upon thermal treatment is indicative of water mobility and consistent with a lack of communication between the pores of MOF layers. For zhemchuzh-

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nikovite, the more extensive loss of water, including that coordinated to Mg^{2+} ions in pores, can be explained by the existence of channels formed by stacking of MOF layers.

The difference in crystal structures of stepanovite and zhemchuzhnikovite points to the possibility of modifying the stacking of **hcb** sheets by forming a solid solution of multiple metal ions (25). Because differences between the structures of zhemchuzhnikovite and stepanovite also lead to different hydrogen-bonded motifs involving guest water molecules, this control over the stacking of MOF sheets would be of potential interest in proton-conducting materials (11). To evaluate how the Al/Fe ratio influences the formation of zhemchuzhnikovite structure, we used mechanochemical liquid-assisted grinding (LAG) (26) to synthesize solid solutions from premade NaMgAl(\mathbf{ox})₃·9H₂O and NaMgFe(\mathbf{ox})₃·9H₂O (stepanovite). LAG was shown to be an efficient method of preparing solid solutions of coordination polymers (27, 28). PXRD analysis of samples prepared by milling NaMgAl(\mathbf{ox})₃·9H₂O and NaMgFe(\mathbf{ox})₃·9H₂O in different stoichiometric ratios for 30 min,

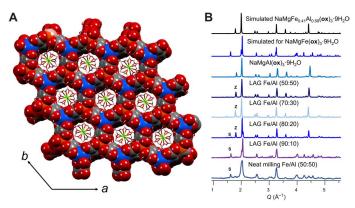


Fig. 3. Crystal structure of zhemchuzhnikovite and role of aluminum in its formation. (A) Zhemchuzhnikovite structure viewed down the crystallographic *c* axis, demonstrating the alignment of MOF layers into channels occupied by $Mg(H_2O)_6^{2+}$. For clarity, the **hcb** layers are shown in spacefilling mode, $Mg(H_2O)_6^{2+}$, and interstitial water molecules using capped sticks. (B) PXRD patterns (top to bottom): simulated NaMgFe_{0.41}Al_{0.59}(**ox**)₃·9H₂O, simulated NaMgFe(**ox**)₃·9H₂O, NaMgAl(**ox**)₃·9H₂O, products of LAG of NaMgFe(**ox**)₃·9H₂O and NaMgAl(**ox**)₃·9H₂O in different stoichiometric ratios (50:50, 70:30, 80:20, and 90:10), and product of dry milling of NaMgFe(**ox**)₃·9H₂O and NaMgAl(**ox**)₃·9H₂O in 50:50 ratio. The PXRD patterns reveal formation of zhemchuzhnikovite structure for Fe/Al ratios up to ~80:20, whereas a higher Fe/Al ratio favors stepanovite structure. Two characteristic reflections that distinguish zhemchuzhnikovite and stepanovite structures are highlighted by "z" and "s," respectively.

using a small amount of water as a grinding liquid (20 µl per 200 mg solid mixture), reveals the zhemchuzhnikovite structure forming with Fe/Al ratios up to 0.8:0.2 (Fig. 3B). At higher ratios, the stepanovite structure becomes preferred. LAG of synthetic stepanovite on its own did not lead to a change in PXRD pattern, indicating that the appearance of the zhemchuzhnikovite structure is tentatively related to NaMgAl(\mathbf{ox})₃·9H₂O, because of the formation of NaMg(Fe_{1-x}Al_x) (ox)₃·9H₂O solid solutions. Fourier transform infrared attenuated total reflectance (ATR) spectra of LAG products exhibit absorption bands resembling NaMgAl(ox)₃·9H₂O and NaMgFe(ox)₃·9H₂O (fig. S11). Dry milling of a 1:1 stoichiometric mixture of NaMgFe(ox)₃·9H₂O and NaMgAl(ox)₃·9H₂O for 30 min results largely in the broadening of x-ray reflections, with stepanovite and zhemchuzhnikovite phases still discernable in the PXRD pattern (Fig. 3B). However, adding water to the dry-milled mixture (20 µl per 200 mg) and milling for five more minutes again produced the zhemchuzhnikovite structure, consistent with LAG.

DISCUSSION

Establishing that open MOF structures can form in a natural environment demonstrates that coordination-driven self-assembly used in creating advanced materials in the laboratory can also play a role in geological processes. Whereas channels and voids are well known for inorganic minerals (29, 30), zhemchuzhnikovite and stepanovite are so far the only examples of organic minerals exhibiting open structures. The structure of zhemchuzhnikovite, with ~0.9-nm channels, is particularly remarkable among organic minerals. Although neither zhemchuzhnikovite nor stepanovite can exhibit permanent porosity, because of the presence of

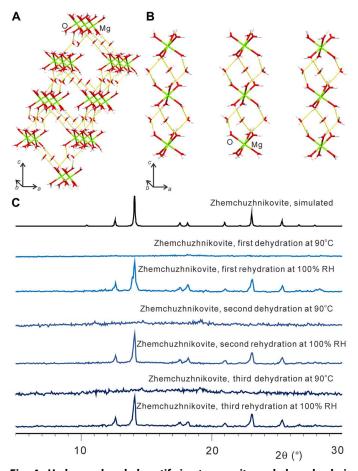


Fig. 4. Hydrogen-bonded motifs in stepanovite and zhemchuzhnikovite and reversibility of thermal dehydration of zhemchuzhnikovite. (A) The 3D hydrogen-bonded network of $Mg(H_2O)_6^{2+}$ ions and interlayer water molecules in stepanovite. (B) Three parallel hydrogenbonded columns of $Mg(H_2O)_6^{2+}$ ions and interlayer water molecules in zhemchuzhnikovite, each propagating through a channel formed by the overlap of hcb layers. For clarity, the metal-organic hcb layers are omitted. (C) The reversibility of structural changes upon dehydration and rehydration of synthetic zhemchuzhnikovite is demonstrated by comparison of PXRD patterns (top to bottom): simulated zhemchuzhnikovite NaMgFe_{0.41}Al_{0.59}(ox)₃·9H₂O; simulated for stepanovite NaMgFe(ox)₃·9H₂O; and zhemchuzhnikovite after one, two, and three cycles of thermal dehydration at 90°C and rehydration by exposure to 100% relative humidity (RH) at room temperature.

 $Mg(H_2O)_6^{2+}$ guests in the MOF pores, they exhibit similarities to previously reported analogous MOFs (10, 11, 19–21). In particular, both minerals can undergo reversible loss and sorption of water guests, demonstrating the stability of individual **hcb** layers upon desolvation, and exhibit extended hydrogen-bonded architectures that suggest the potential for proton conduction (24). The open metal-organic architectures in zhemchuzhnikovite and stepanovite change our view of MOFs as strictly artificial materials and hint to the possibility that the future may unravel other MOF minerals, potentially even microporous ones. These expectations are not far-fetched, bearing in mind that some organic minerals involve multifunctional molecules already used in MOF synthesis, such as mellitic acid or different purines (12, 31–34).

MATERIALS AND METHODS

Solution synthesis

NaMgAl(**ox**)₃·9H₂O was prepared according to a published procedure (22). Synthetic stepanovite was prepared by dissolving stoichiometric amounts of NaOH, MgO (\geq 99% trace metals basis, ~325-mesh powder; Sigma-Aldrich), and Fe₂O₃ (nanopowder; Sigma-Aldrich) in aqueous oxalic acid solution. The resulting green solution was stirred overnight, filtered, and left to evaporate at 4°C. Green, rhombohedral crystals were obtained after 2 days. Synthetic zhemchuzhnikovite was obtained by the evaporation of an aqueous solution obtained by mixing saturated aqueous Na₃Fe(**ox**)₃·3H₂O, aluminum sulfate, and MgO with aqueous oxalic acid, using a 6:4 Al/Fe stoichiometric ratio. After 2 days at 4°C, yellow-green needles formed.

Single-crystal XRD

Crystal structures of synthetic and natural stepanovite and zhemchuzhnikovite were determined by single-crystal XRD. Diffraction measurements were made on Bruker D8 APEX2 and Bruker APEX DUO x-ray diffractometers, using graphite-monochromated Mo K_{α} radiation ($\lambda = 0.71073$ Å). Data were collected in ω scan mode ($2\theta \leq$ 54°). Structures were solved by direct methods in SHELXS and refined using SHELXL (35) on F^2 using all data. Hydrogen atoms were located using the electron difference map when permitted by data quality. Calculations were performed and images were prepared using WinGX program suite (36). Structures have been deposited to the CSD, with deposition codes 1408093 to 1408095 for synthetic samples and 1431678 and 1431679 for natural samples.

Mechanochemical milling

Mechanochemical milling was performed for 30 min, using an MM200 Retsch mill operating at 25 Hz, in custom-made Teflon jars (14-ml volume), and with two stainless steel balls (each ball was 8 mm in diameter and 1.9 g in weight).

Powder x-ray diffraction

Powder x-ray diffraction data for synthetic samples were collected on Bruker D2 LYNXEYE Phaser using nickel-filtered Cu K_{α} radiation ($\lambda = 1.54056$ Å), and those for natural samples were collected using a Rigaku R-AXIS Rapid II diffractometer (curved image plate, d = 127.4 mm; Debye-Scherrer geometry; Co K_{α} , $\lambda = 1.79021$ Å).

Cambridge Structural Database

CSD searches were performed using CSD version 5.36 (November 2014) +1 update.

Fourier transform infrared ATR

Fourier transform infrared ATR was measured using Bruker VER-TEX 70 with PLATINUM diamond ATR unit.

Thermogravimetric analysis

Thermogravimetric analysis was performed using a Mettler-Toledo TGA DSC 1 Star thermobalance (using alumina crucibles) and TA Instruments Q500 (using platinum pans) in a stream of air (flow rate, 60 ml min⁻¹; heating rate, 10°C min⁻¹; sample size, 5 to 10 mg). Data were processed using Mettler STARe 9.01 Software.

SUPPLEMENTARY MATERIALS

Supplementary material for this article is available at http://advances.sciencemag.org/cgi/ content/full/2/8/e1600621/DC1

fig. S1. Scanning electron microscopy image of aggregates of isometric zhemchuzhnikovite crystals from Chai-Tumus coal deposit, Sakha-Yakutia, Russia (sample from E. I. Nefedov's collections).

fig. S2. Morphologies for stepanovite, reported by Knipovich et al. (8).

- fig. S3. Morphologies for zhemchuzhnikovite.
- fig. S4. TGA of bulk synthetic stepanovite.
- fig. S5. TGA of bulk synthetic zhemchuzhnikovite.

fig. S6. Comparison of PXRD patterns for synthetic stepanovite after one, two, and three cycles of dehydration at 90°C and rehydration at 100% RH.

fig. S7. Thermal analysis of zhemchuzhnikovite after thermal dehydration at 90°C for 16 hours: TGA (top) and DSC (bottom).

fig. S8. Thermal analysis of zhemchuzhnikovite after thermal dehydration at 90°C, followed by exposure to 100% RH: TGA (top) and DSC (bottom).

fig. S9. Thermal analysis of stepanovite after thermal dehydration at 90°C: TGA (top) and DSC (bottom). fig. S10. Thermal analysis of stepanovite after thermal dehydration at 90°C, followed by exposure to 100% RH: TGA (top) and DSC (bottom).

fig. S11. Overlay of FTIR-ATR spectra for mechanochemically prepared zhemchuzhnikovite analogs with different AI:Fe ratios.

table S1. PXRD data of bulk stepanovite from Chai-Tumus coal deposit, Siberia, Russia, were indexed on the basis of herein determined crystal structure.

table S2. PXRD data of bulk zhemchuzhnikovite from Chai-Tumus coal deposit, Siberia, Russia, were indexed on the basis of herein determined crystal structure.

data file S1. Crystallographic data for crystal structures of natural and synthetic samples of stepanovite and zhemchuzhnikovite.

data file S2. checkCIF for crystal structures of natural and synthetic samples of stepanovite and zhemchuzhnikovite.

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